



Laboratoire de
Géologie

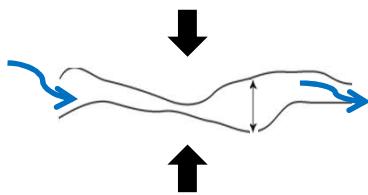
Permeability of cracked rocks and glass

Yves GUEGUEN,
Jérôme FORTIN
Audrey OUGIER-SIMONIN



Permeability in the upper crust

- Basaltic ocean crust permeability decrease from near surface values (100 mD) to low values (< 0.1 mD) at 1km and deeper (Fisher, 1998)
- From log data, evidence of crack networks (Violay, 2010)
- Permeability-depth variation can be interpreted by crack closure (low T conditions, no sealing)



Cracks play a major role



Low permeability rocks under confining pressure

Reference	Rock type (low porosity rocks)	a, 10^{-2} MPa $^{-1}$	P* = a $^{-1}$, MPa
YALE (1984)	tight sandstones	3.8 - 6.3	15.9 - 26.3
BRACE <i>et al.</i> (1968)	Westerly granite	3.3	30.6
BERNABE (1986)	Chelmsford granite	2.9	34.6
	Barre granite	2.3	42.7
FISHER and PATERSON (1992)	Carrara marble	4.7	21.3
HUENGES and WILL (1989)	Urach gneiss	4.9	20.4
MORROW <i>et al.</i> (1994)	gneiss [Kola]	3.2	31.2
	basalt [Kola]	10.2	9.8
	amphibolite [KTB]	5.8 -11	9.1 - 17.2

$$k = k_o e^{-aP}$$



Cracks ?

10 < P* < 40
(MPa)



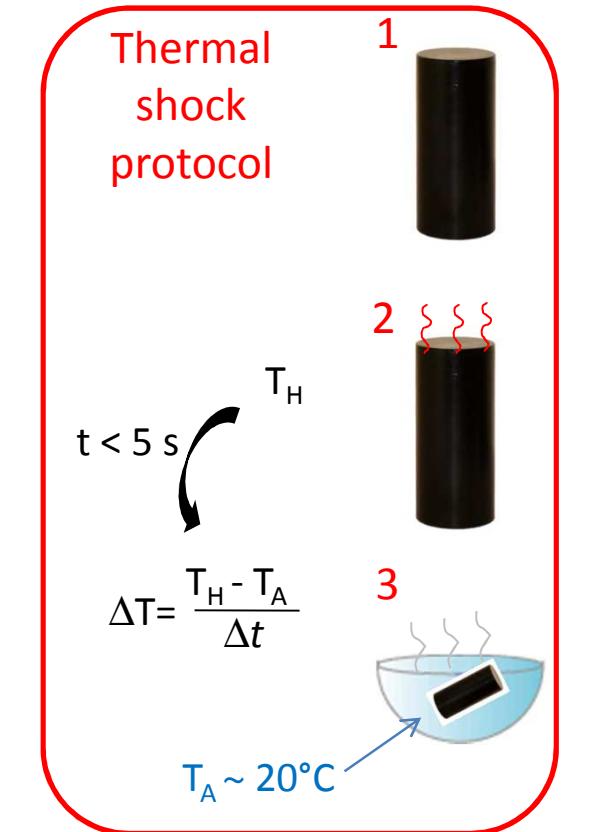
An absolute reference: glass + cracks

Low permeability rocks : cracks ?

An ideal case: intact synthetic glass with cracks

How to introduce cracks in glass samples ?

- thermal shock
- cracks only

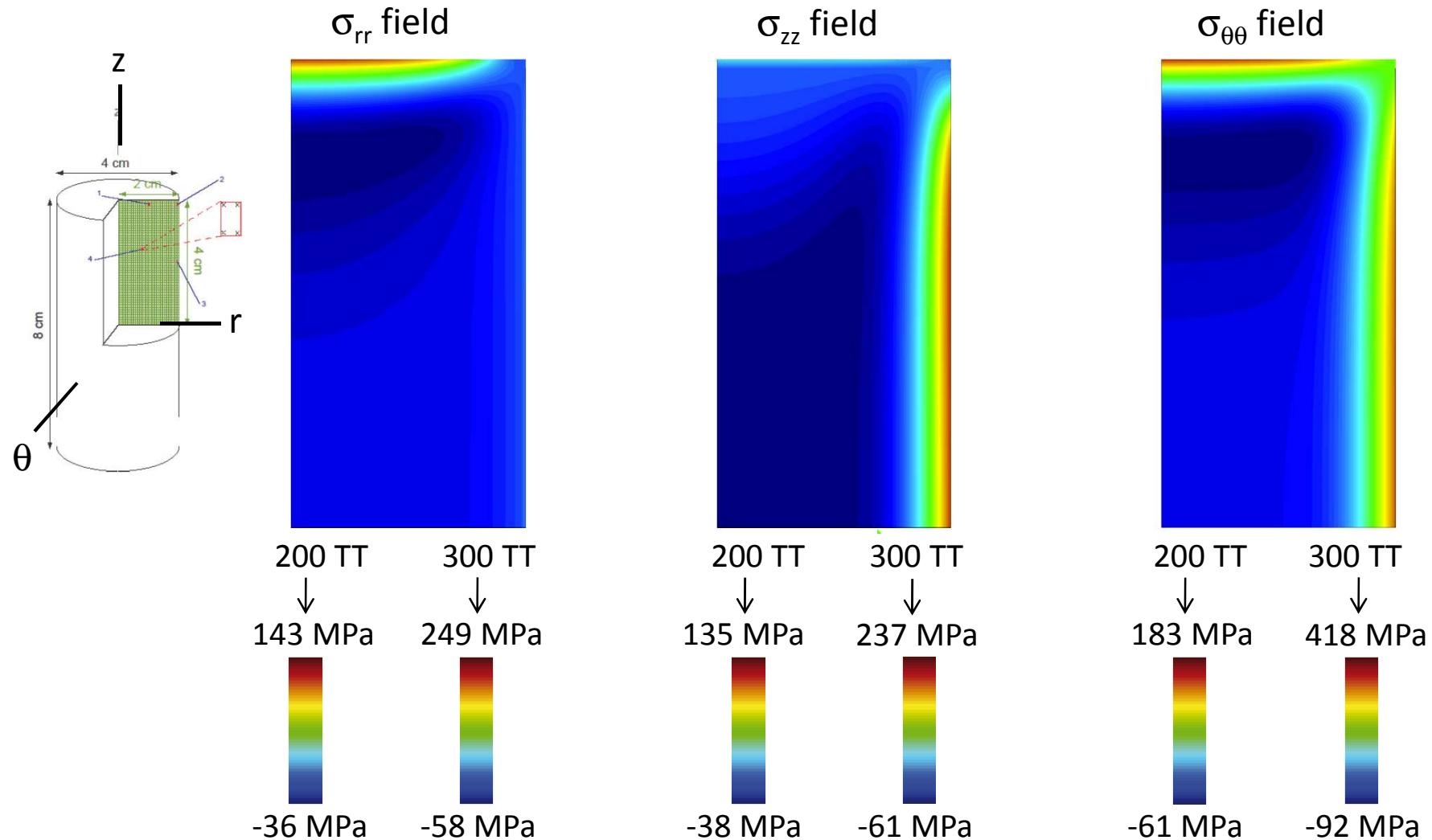


$$T_H = 100, 200 \text{ and } 300^\circ\text{C}$$



Shocked glasses: temperature and stress distribution

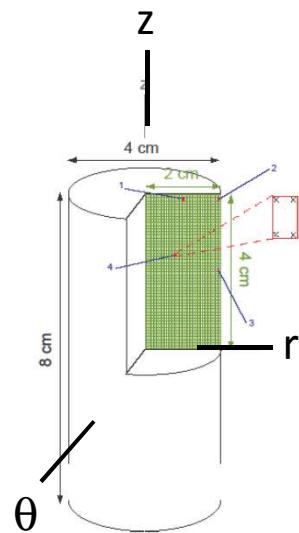
$t = 20\text{s}$ after thermal shock





Shocked glasses: temperature and stress distribution

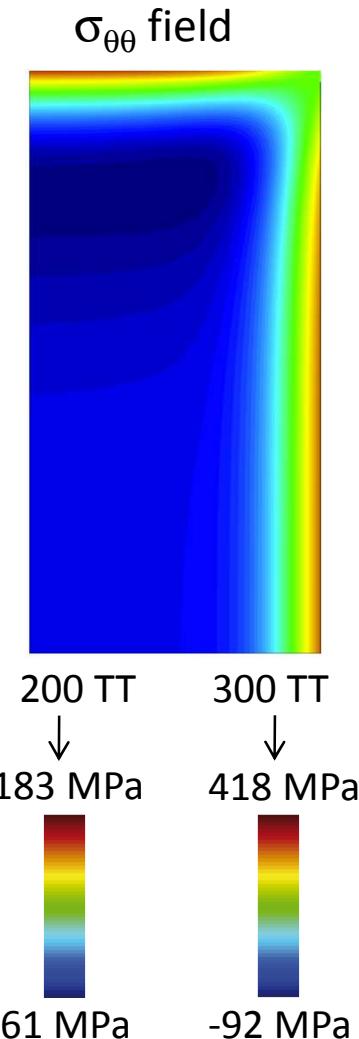
$t = 20\text{s}$ after thermal shock

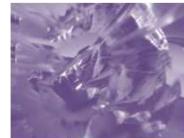


Outward ring : in traction
Inner core: in compression

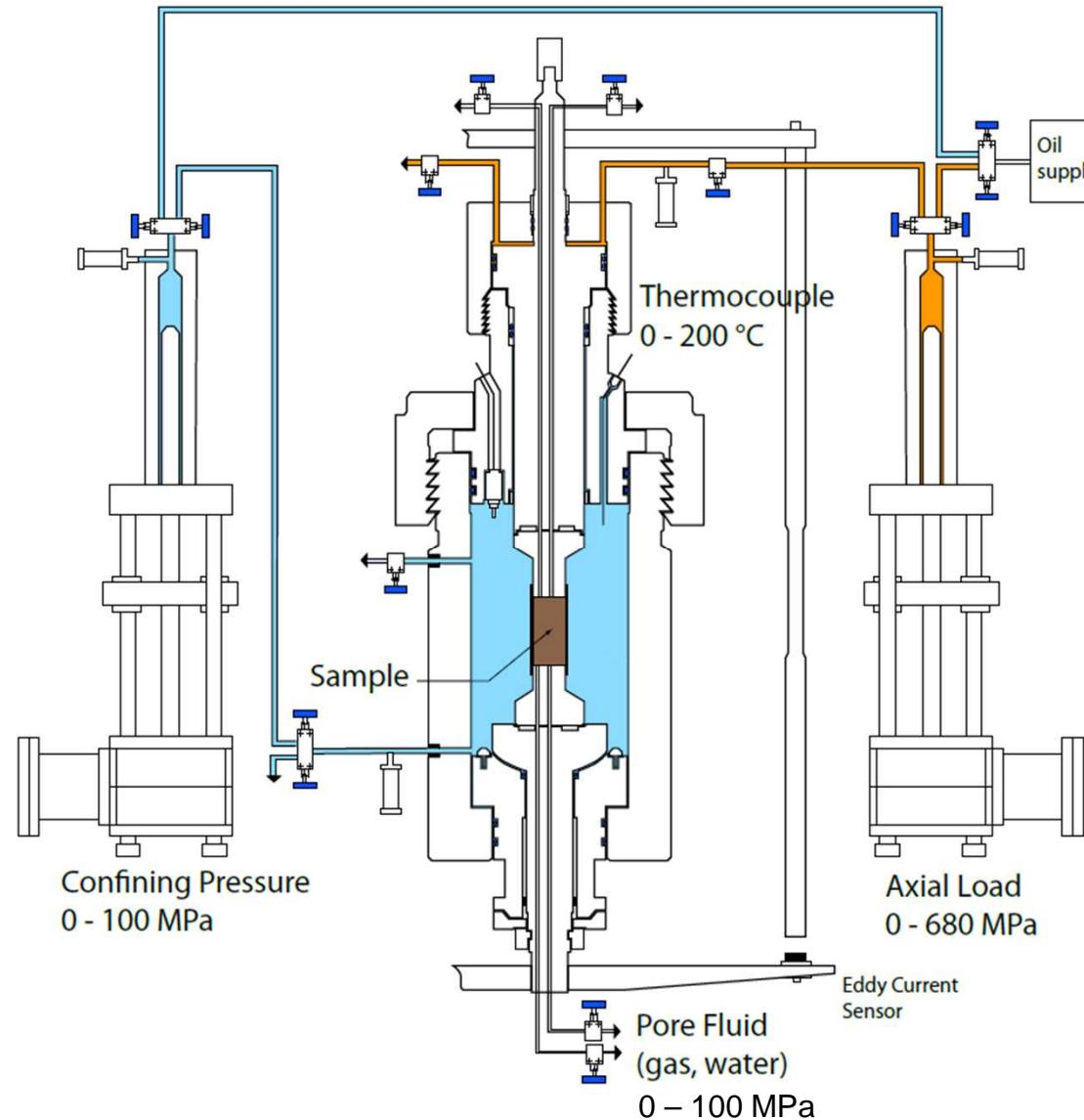


Mostly vertical cracks in the outward ring





Triaxial cell apparatus: Measurement of V_p , V_s , k



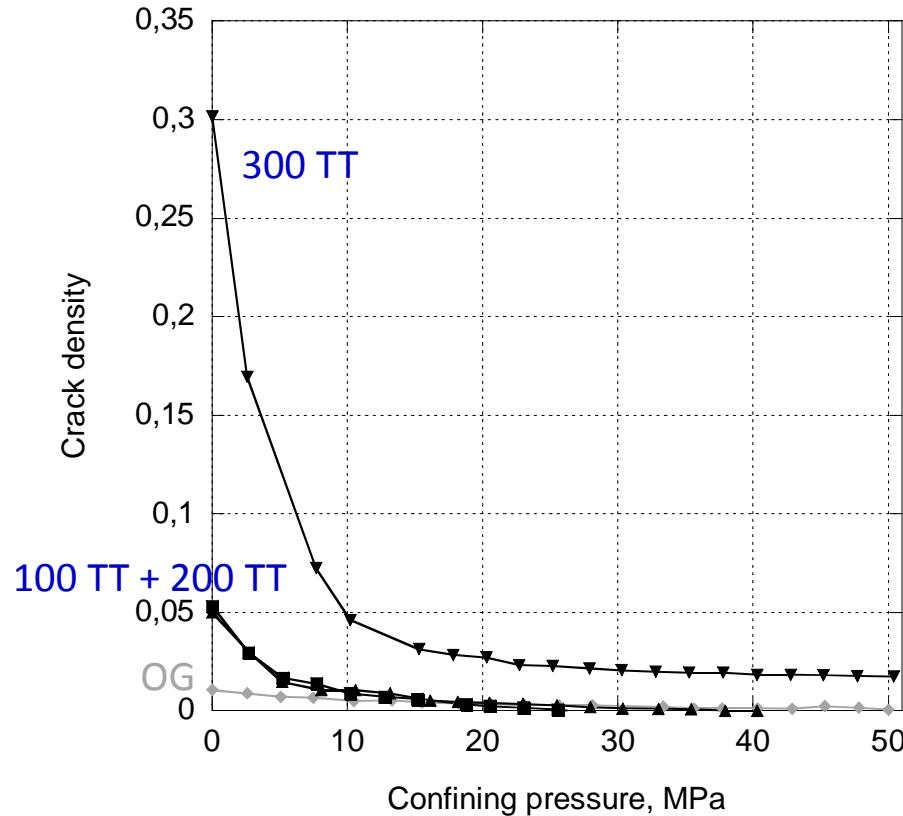


Crack density (from Vp, Vs) variation with pressure

Crack density is obtained from elastic moduli K and G (Vp and Vs)

$$\left\{ \begin{array}{l} \frac{K_o}{K} = 1 + \rho_c \frac{h}{1 - 2\nu_o} \left(1 - \frac{\nu_o}{2}\right) \\ \frac{G_o}{G} = 1 + \rho_c \frac{h}{1 + \nu_o} \left(1 - \frac{\nu_o}{5}\right) \end{array} \right.$$

Crack aspect ratio ξ



→ Crack closure:

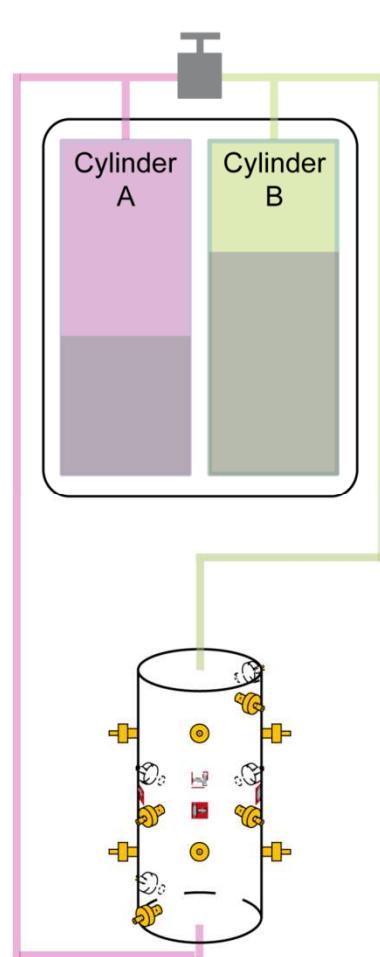
at $P_{\text{hydro}} \approx 15 \text{ MPa}$ for shocks at 100 and 200 °C

at $P_{\text{hydro}} \approx 40 \text{ MPa}$ for shock at 300 °C

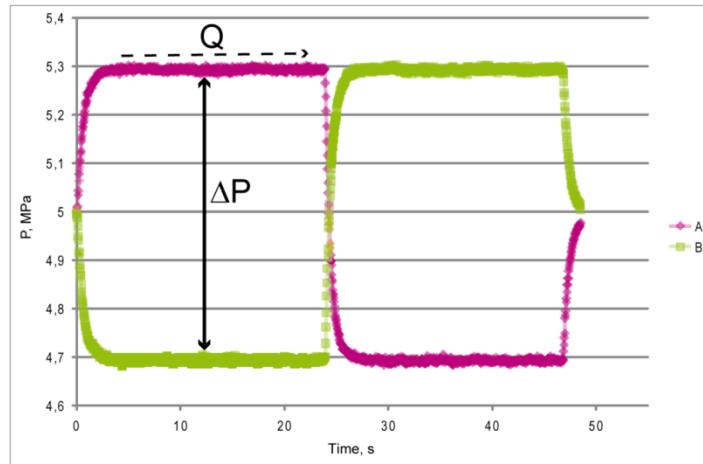
→ $\xi \sim 10^{-3} \text{ to } 10^{-4}$



Methods of measurements of permeability



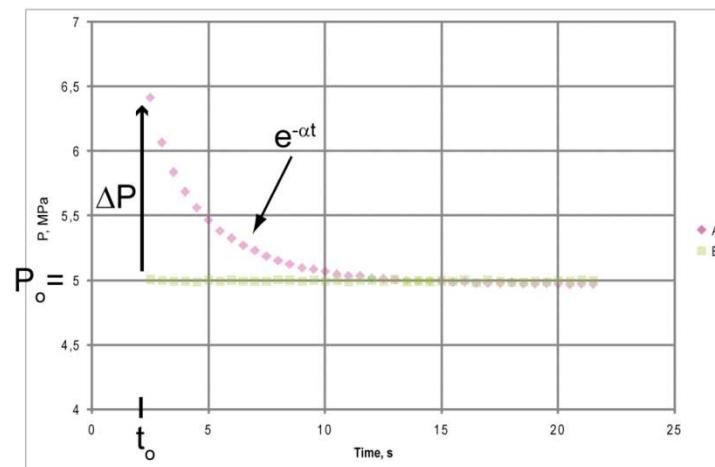
a) Constant flow method



→ $P_{\text{eff}} < 10 \text{ MPa}$

$$\text{Darcy's law: } k = \frac{Q}{\Delta P} \frac{L}{A} \mu$$

b) Pulse decay method (1/2-pulse)



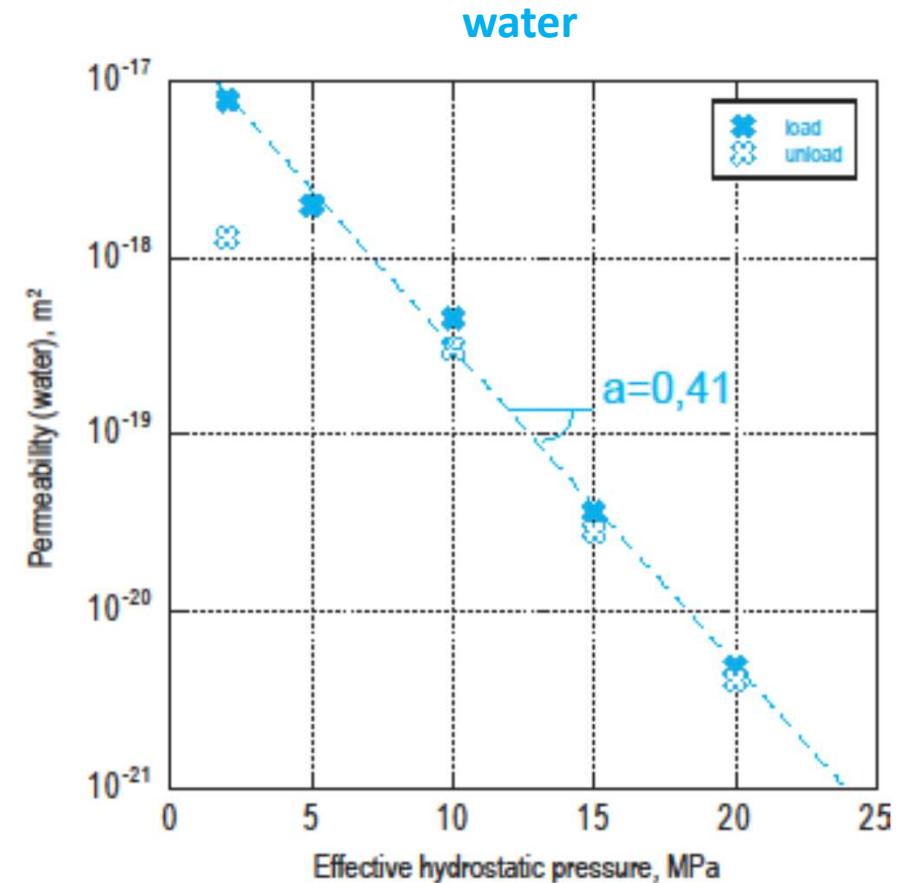
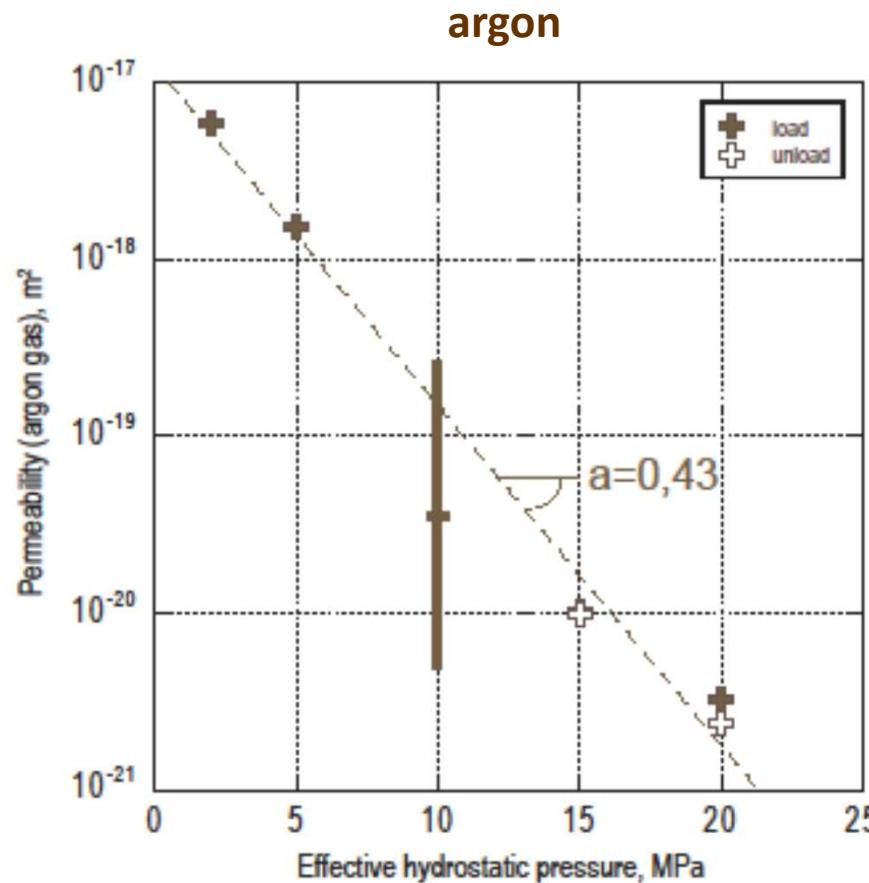
→ $P_{\text{effective}} \geq 10 \text{ MPa}$

$$k = \alpha \beta \frac{L}{A} \mu V_{\text{pulse}}$$

Brace et al. (1968)



Cracked glass permeability under confining pressure



$$k = k_o e^{-aP} \quad \text{and} \quad a^{-1} = P^* \sim 2.4 \text{ MPa}$$



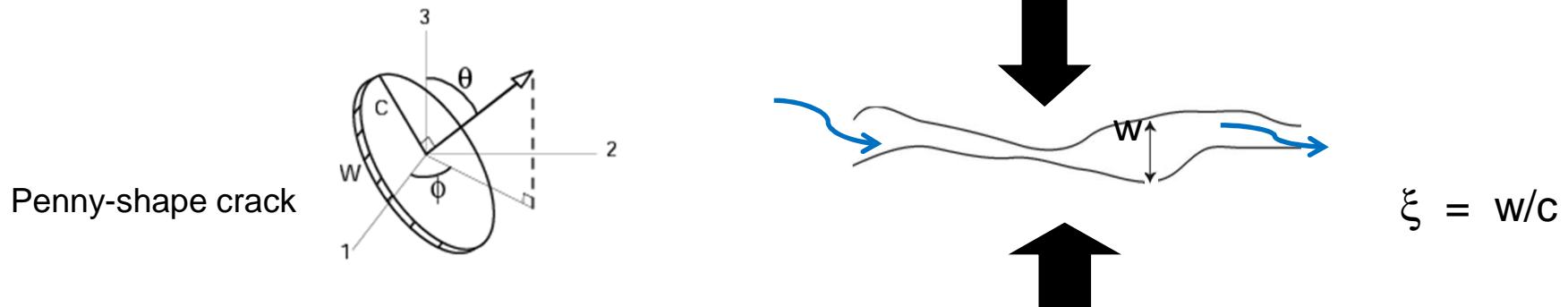
Comparison with low permeability rocks

Reference	Rock type (low porosity rocks)	$a, 10^{-2} \text{ MPa}^{-1}$	$P^* = a^{-1}, \text{ MPa}$
YALE (1984)	tight sandstones	3.8 - 6.3	15.9 - 26.3
BRACE <i>et al.</i> (1968)	Westerly granite	3.3	30.6
BERNABE (1986)	Chelmsford granite	2.9	34.6
	Barre granite	2.3	42.7
FISHER and PATERSON (1992)	Carrara marble	4.7	21.3
HUENGES and WILL (1989)	Urach gneiss	4.9	20.4
MORROW <i>et al.</i> (1994)	gneiss [Kola]	3.2	31.2
	basalt [Kola]	10.2	9.8
	amphibolite [KTB]	5.8 - 11	9.1 - 17.2
SON68 glass		$a = 0.4 \text{ MPa}^{-1}$	$P^* = 2.4 \text{ MPa}$

$$k = k_o e^{-aP}$$



Permeability and crack closure



data $\rightarrow k = k_o e^{-aP}$

Crack elastic closure $\rightarrow w \sim w_o \left(1 - \frac{P}{E\xi}\right)$

Cubic law $\rightarrow k \sim w^3$



P^* value in cracked rocks and glass: $k = k_0 e^{-P/P^*}$



$$k = k_o e^{-aP}$$

$$k_o e^{-aP} \propto w_o^3 \left(1 - \frac{P}{E\xi}\right)^3$$

$$w \sim w_o \left(1 - \frac{P}{E\xi}\right)$$

This implies :

$$a = \frac{3}{E\xi}$$

or :

$$P^* = E\xi / 3$$



$$P^* \text{ value in glass : } k = k_0 e^{-P/P^*}$$



$$P^* = E\xi / 3$$

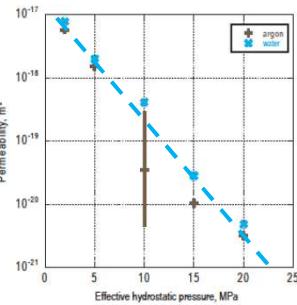
$$\text{With : } E = 80 \text{ GPa} \quad \text{and} \quad \xi = 10^{-4} \quad \longrightarrow \quad P^* = 2.6 \text{ MPa}$$

(Measured value: $P^* = 2.4 \text{ MPa}$)

In glass, no microstructure (but a nanostructure) \longrightarrow Very sharp crack tip

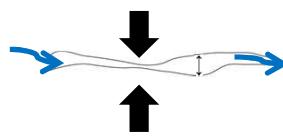


Conclusions: permeability of cracked rocks and glass



- Permeability decreases follows an exponential law in cracked rocks and glass

- $k = k_0 e^{-P/P^*}$, where $P^* = E\xi / 3$



- Very sharp crack tips in glass : $P^* = 2 - 3 \text{ MPa}$, aspect ratio

$$\xi \sim 10^{-4}$$

- In rocks, $P^* = 10 - 40 \text{ MPa}$

aspect ratio > 10^{-3}



Further work in progress (see C. Mallet poster)