Creep modelling of soft soils: September 2014

Introduction

From field and laboratory observations to simple creep models and their parameters Towards extended creep models

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(Dundee University, Bristol University)







Creep modelling of clay - stress/strain reversal



– p. 1/3

kaolin revisited: yield points? one-dimensional consolidation histories



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data from Al-Tabbaa (1987)



plastic strain increments: approximate normality to kinematic yield loci





kinematic hardening extension

yield locus carried around with stress state – 'bubble' – strongly influenced by recent history

stiffness falls as yield 'bubble' approaches bounding surface – controlled by distance *b*

when loading with 'bubble' in contact with bounding surface model is identical to Cam clay



creep modelling of clay - stress reversal



creep modelling of clay - stress reversal





creep modelling of clay - stress reversal

- reversal of stress or strain path for clays (and other soils) shows reverse plasticity and hysteretic, dissipative response in unloading-reloading cycles
- yielding of clays (and other soils) is convincingly described using kinematic hardening combined with bounding surface plasticity
- anticipate that viscoplasticity of clays should also be described by a kinematic hardening/bounding surface modelling framework
- 'overstress' approach applied to two mechanisms of creep: referred to 'bounding surface' and to kinematic yield surface ('bubble')
- expect 'loading' creep to dominate as unloading begins
- expect 'reversal' creep to dominate as overconsolidation increases

Creep modelling of peat = soil + fibres?



- p. 2/3

7. Applications?



- Peat: partially decayed plant matter
- recognisable mass of roots and woody tissue at the surface \Rightarrow partially decayed layer \Rightarrow soil-like material at depth of ${\sim}10m$
- highly compressible
- important contribution of fibrous material



4. Mixtures of soil (sand) with flexible fibres



flexible polypropylene fibres

fibres 35mm long, 0.1mm diameter

analogy for roots?

mixtures with Hostun sand $d_{50} = 0.38$ mm, $C_u = 1.9$



4. Mixtures of soil (sand) with flexible fibres



deduce distribution of orientations

moist tamping leaves most fibres within $\pi/4$ of horizontal



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hypothesis 1: tensile strains in soil try to stretch fibres





 hypothesis 2: stretched fibres tend to increase normal stress on soil and contribute to shear stress



tensile strain in soil ...



... not completely transferred to fibre slip depends on normal stress

- hypothesis 3: bond between fibres and soil not perfect
- $\varepsilon_f / \varepsilon_m = 1 \lambda \exp(\sigma'_z / \sigma_{rf})$
- ε_f strain in fibre; ε_m strain in soil







 hypothesis 4: fibres (may pull out of soil or) may reach tensile strength and snap



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 hypothesis 4: fibres may pull out of soil or may reach tensile strength and snap





unit area of cross-section

 hypothesis 5: fibres treated as forces with orientation (not continuous material)





 Severn-Trent sand: simple mathematical relationships: automatic convergence on critical state

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- simulations of shear box tests on fibre-sand mixtures
- fibre orientations $-\pi/4 < \theta < \pi/4$
- dominant effect from fibrespace densification



– p. 2/1

7. Applications?



- Peat: partially decayed plant matter
- recognisable mass of roots and woody tissue at the surface \Rightarrow partially decayed layer \Rightarrow soil-like material at depth of ${\sim}10m$
- highly compressible
- important contribution of fibrous material

treat as fibres + clay?





 hypothesis 2: stretched fibres tend to increase normal stress on soil and contribute to shear stress viscous effects in soil:fibre interaction?
tensile strain in soil ...



... not completely transferred to fibre slip depends on normal stress

- hypothesis 3: bond between fibres and soil not perfect
- $\varepsilon_f / \varepsilon_m = 1 \lambda \exp(\sigma'_z / \sigma_{rf})$
- ε_f strain in fibre; ε_m strain in soil



viscous effects in soil:fibre interaction?



 hypothesis 4: fibres may pull out of soil or may reach tensile strength and snap





- simulations of shear box tests on fibre-sand mixtures
- fibre orientations $-\pi/4 < \theta < \pi/4$
- dominant effect from fibrespace densification

fibre-clay?

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Creep modelling of sand - role of particle breakage/damage



creep modelling of sand





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(a) (b)

Figure 1. SEM images of Ottawa sand and Lake Michigan Dune Sand grain surfaces (image width $\approx 100 \,\mu$ m).

Figure 4. Damage of a large asperity at a contact loaded with 0.7 N (SEM image width $70 \approx \mu \text{m}$).



Figure 2. SEM image of Lake Michigan Dune Sand (image width $\approx 10 \,\mu$ m).



Figure 3. AFM scan of Ottawa sand grain surface $(2 \times 2 \,\mu\text{m})$.

particle breakage/suffusion



- particle breakage broadens grading
- internal erosion (suffusion) narrows grading



particle breakage/suffusion



- broader gradings pack more efficiently
- limiting densities increase
- critical state line falls
- state parameter increases soil feels looser



particle breakage/suffusion



- characterise grading and grading evolution
- link grading evolution and critical state line
- particle breakage criterion stress (strain?)
- other aspects of model unchanged (first order)



particle breakage + Severn-Trent sand





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particle breakage + Severn-Trent sand





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creep modelling of sands

- complex patterns of response observed 'descriptive' titles but little attempt to understand physical mechanisms (Tatsuoka *et al.*)
- time effects in sands relate to the amount of grain crushing (Lade)
- creep effects observed in absence of particle breakage (Airey)
- experimental evidence consistent with static fatigue: time dependent fracturing of micro-morphologic features at inter-granular contacts (Michalowski)
- model for soil with varying particle size distribution


Creep of Soft Clay: "Classical" to "Current Practice"

Hans Petter Jostad

Discipline leader in numerical modeling at NGI Adjunct Professor at NTNU

2st CREEP Course, Trondheim, Norway,

15-16. September 2014





Key questions:

What is the definition of creep?

Do we have creep deformations at the same time as we have deformations due to pore pressure/stress changes?

- When does creep start?
- What controls the (volumetric) creep deformation?

How to expand from 1D to a general 3D stress state?

Motivation

How to calculate long term settlements in soft clay? A) Primary and secondary compression phases? B) Coupled consolidation and creep?

$$\frac{de}{dt} = \left(\frac{\partial e}{\partial \sigma'}\right)_{t} \frac{d\sigma'}{dt} + \left(\frac{\partial e}{\partial t}\right)_{\sigma'} + \left\{\frac{\partial e}{\partial T} \frac{dT}{dt}\right\} + \left\{\frac{\partial e}{\partial \theta_{i}} \frac{d\theta_{i}}{dt}\right\}$$
stress induced creep (temperature) (chemical)
$$\Delta e = \int_{0}^{t_{p}} \left[\left(\frac{\partial e}{\partial \sigma'}\right)_{t} \frac{d\sigma'}{dt} + \left(\frac{\partial e}{\partial t}\right)_{\sigma'} \right] dt + \int_{t_{p}}^{t} \left[\left(\frac{\partial e}{\partial t}\right)_{\sigma'} \right] dt$$
Primary (consolidation) Secondary (creep)

Secondary consolidation (creep)



The classical approach for creep settlements!



Bjerrum's delayed compression concept



VERTICAL PRESSURE IN LOGARITHMIC SCALE

Unique relationship between creep rate, effective (vertical) stress and void ratio



Key questions related to long term settlements

Extrapolation of laboratory data; rate?



Behaviour around pc' (destructuration)?



Extrapolation of laboratory data; time?



Effect of varying load history?



Unique end-of-primary (EOP) void ratio



Need this assumption in order to divide into primary- and secondary phases A or B most correct?

Unique end-of-primary (EOP) void ratio

Arguments against the approach:

- How does a soil element "feel" that it is in a primary consolidation state?
 - Elements close to a drainage condition with almost constant effective stresses during the consolidation phase
 - Time of primary consolidation governed by a low permeable layer
- Difficult to define a unique EOP state
 - $u_{excess}/q = 0.1, 0.05 \text{ or } 0.01$

Main challenges

Due to significantly different time scale in field and laboratory conditions, the deformation in the field must be described (extrapolated) by a creep model based on input from laboratory tests.



Sample disturbance (apparent creep)



Is soil disturbance the reason for good agreements between standard consolidation analyses (without creep) and field observation?

 $M_{oc} = a \cdot M_{pc}$ a = 5 -> 15



Verification/calibration of creep models

Back-analyses of measured field data Generally large number of uncertainties
Back-analyses of idealised model tests (e.g. oedometer tests with different specimen heights)
Extrapolation is still necessary
Long term laboratory tests
Extrapolation is still necessary

Oedometer test (Incremental Loading Tests)



Need to separate into contributions from effective stress changes and creep:

- Creep "starts" when u ~ 0 or after 1 day (as a reference)?
- Or, all plastic strains are time dependent (Soft Soil Creep Model)

Janbu's resistance concept (EP+VP)



Creep is added to the elasto-plastic strains



Janbu's time resistance

$$R = R_o + r (t - t_o)$$

$$\Delta \varepsilon_{creep} = \frac{1}{r} \ln \left(\frac{t}{t_o} \right) = \frac{1}{r} \ln \left(\frac{R}{R_o} \right) \longrightarrow t_{eqv}$$

$$\Delta \varepsilon_{creep} = \varepsilon_{tot} - \varepsilon_o(\sigma_v')$$

 $R = R_o e^{\left(r \, \Delta \varepsilon_{creep}\right)}$

The accumulate creep strain is the state parameter for creep rate



Effective stress dependency

Stress or OCR dependent?









Ro = 0.3 and 1 year, r = 100, 300 and 500



A large contribution of creep may occur during primary consolidation



Isotaches – lines of constant (creep) strain rate

unique relationship between effective stress – strain (void ratio) – strain rate



These curves may be non-linear (curved)!



"Apparent" pre-consolidation pressure Plaxis - Soft Soil Creep Model (E+EVP)



$$\Delta \varepsilon_{creep} = \frac{1}{r} \ln \left(\frac{R}{R_o} \right) = \mu^* \ln \left(\frac{R}{R_o} \right) = (\lambda^* - \kappa^*) \ln \left(\frac{\sigma_{vc}}{\sigma_{v'}} \right) = (\lambda^* - \kappa^*) \ln \left(OCR \right)$$

$$R = R_o e^{\left(r \,\Delta \varepsilon_{creep}\right)} = R_o OCR^{\frac{\lambda^* - \kappa^*}{\mu^*}} = \frac{t_{ref}}{\mu^*} OCR^{\frac{\lambda^* - \kappa^*}{\mu^*}}$$



Solution algorithm – FE program



Soil Investigation

Soil profile from e.g. CPTU and location of depth to bedrock (or a stiff layer)

In-situ pore pressure measurements (piezometers)

Soil samples from different depths/soil layers

Standard index tests

Oedometer tests

- constant strain rate (CRS) tests with unloading/reloading loops.
- x days creep test (and/or CRS tests with different rates)
- additional permeability tests?
- incremental loading (IL) tests (specification: Dq/q=1?, duration=24 hours or EOP, pore pressure measurements, long term creep phases, etc)



Recommendations of laboratory tests

IL tests are well suited to provide data on creep parameters and the location of the RTL

CRS tests is recommended to define the shape of one isotache specially around the yield stress

How should we define creep behaviour before pc?





Peat / organic soils

Dr ir C. Zwanenburg

23 september 2014

<u>contents</u>

- Intro
- Engineering problems
- Characterisation
- Engineering parameters
- Settlement and creep
- shortcomings

What is peat / organic soil?

Deposit of organic material, to some extent mixed with clay, silt sand particles





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(http://www.grida.no/graphicslib/detail/peat-distribution-in-the-world_8660)

5-8% total land surface 8-11 % tropical / subtropical (Mesri & Aljouni, 2007)

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Distribution of peat (Nieuwkoop formation) and population density





ENGINEERING PROBLEMS



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Long term settlement





Dewatering



Consequences

adjusting water table leads to extra settlement (location: Waterland)



Case Rotterdam, schiewijk





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Residual settlement





Skin friction on piled foundations



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Case Gouda Goverwelle





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differential settlement





CHARACTERISATION



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Organic content $P = M_{org} / M_{tot}$ Ash content = 1-P

How to determine *P*:

- Determine dry solid mass (M_s), by drying sample for 24 h at 105°C
- Determine remaining mass (m₁) after drying for 4 h at 500 °C
- Loss on ignition $N = (M_s m_1)/M_{s.}$
- Ash content = 1.04(1 *N*)
- P = 1 1.04(1 N)

An error of 4% is assumed for organic particles that are lost in finding M_s

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Definition

	OSRC System		Jarrett System		Davis (1946)	USSR System		L C Sys	GS tem	Landva <i>et</i> <i>al.</i> (1983)
X	Peat	Low Ash Medium Ash High Ash	Pe	at	Peat		1 2 3 4	Pe	at	Peat
A A A A A A A A A A A A A A A A A A A	Carbonaceous Sediment	Low Ash		Peaty		Peat	5	- Peaty	Muck	Peaty Organic Soils
The second s		High Ash	Muck	Clayey / Silty / Sandy / Gravelly	Muck		M1	ick	Organic Soils	
and a second sec	Mineral Sediment		Organic Clay or Silt		Mineral Soil	Non-Peat		Clayey Muck Clay Mucky		

Figure 2.5 Comparison of classification systems used for peat and organic soils (after Andrejko et al., 1983). (B.B.K. Huat et al 2014)

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Characterisation



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Figure 13. Correlation of wet and dry bulk density with natural water content for various Dutch peats and organic soils.

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Further classification

Many options:

- Von Post classification (decomposition)
- Fiberosity
- Botanical background
- Conditions during deposition (eutrophic, mesotrophic, oligotrophic)
- Type of additive (clay, sand etc.)
-

Von Post classification (field identification)

Table A1.Degree of humification (von Post system).

Degree of humification	Decomposition	Plant structure	Content of amorphous material	Material extruded on squeezing (passing between fingers)	Nature of residue
H ₁	None	Easily identified	None	Clear, colourless water	
H_2	Insignificant	Easily identified	None	Yellowish water	
H ₃	Very slight	Still identifiable	Slight	Brown, muddy water; no peat	Not pasty
H ₄	Slight	Not easily identified	Some	Dark brown, muddy water: no peat	Somewhat pasty
H5	Moderate	Recognizable,	Considerable	Muddy water and some peat	Strongly pasty
H ₆	Moderately stong	Indistinct (more distinct after squeezing)	Considerable	About one third of peat squeezed out; water dark brown	
H ₇	Strong	Faintly recognizable	High	About one half of peat squeezed out; any water very dark brown	
H ₈	Very strong	Very indistinct	High	About two thirds of peat squeezed out; also some pasty water	Plant tissue capable of resisting decomposition (roots_fibres)
H9	Nearly complete	Almost not recognizable		Nearly all the peat squeezed out as a fairly uniform paste	(10003, 110103)
H ₁₀	Complete	Not discernible		All the peat passes between the fingers; no free water visible.	(Landva 20

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Fibrosity





50%

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30%

40%



(Source: TR-GCV , TAW 1996)

Either, dry, sieve and count or visual inspection

Botanical background, sedge



(Meier-Uhlherr et al 2011)

Photo G. Erkens

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Botanical background, Sphagnum





Photo G. Erkens 2012









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(Meier-Uhlherr et al 2011)

Conditions during deposition



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(after Visscher 1949, Lowe & Walker 1997)



Example, eutrophic lake



Examples of Marshes and Swamps



Marsh (= low wetland area, covered by reed, grass etc.) Swamp (= low wetland area, covered by forest)

http://sts.gsc.nrcan.gc.ca/



Fen (= low area covered by grass and reed)



Examples of raised bogs



http://sts.gsc.nrcan.gc.ca/







ENGINEERING PARAMETERS



Uitdam test side Map Sati N518 Landsmeer E22 Twiske N247 Rijperweg Zuidsloet Ω Holysloter Sportpark Melkweg Uitdammer Die Kadoelen Tuindorp Oostzaan Zijkaneal I Banne Buiksloot Zunderdorp Bulksloterbeek Marcurlushaven (\$116) Wilhesloot nuweSloot Johan van Hassettkanaal-West Nieuwendammerdijk en Buiksloterdijk Waterland Amsterdam-Noord Kinselmeer Westerpark Haarlemmerbuurt 116 Zijkanaal K naar Nieuwendam n er Jordaan Dijksgracht Zeeburg - 1 Amsterdam Grachtengordel-Zuid Krommerdt aarsjes Nieuwe Diep . Oosterparkbuurt Flevopark rtoomse (E) IJburg West Sluis Van Gogh Groene

Diaman and Junnel

Museum

Characterisation

biological background peat: mainly sedge – reed

Von Post classification: H2 -H3, meaning that the peat is undecomposed (H2) or very slightly decomposed (H3). Plant remains are identifiable and no amorphous material is present.



Profiel proefperceel Uitdam

Locatie profiel in proefperceel:





Erkens, van Wirdum, Wiersma Deltares 2012

Characterisation of peat layer



N [%]



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IJkdijk Stabiliteitsexperiment

Characterisation of peat layer





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IJkdijk Stabiliteitsexperiment

-8

CPTU, 10 cm² cone





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	μ _{qt} , μ _{ball} [MPa]	V _{qt} , V _{qball} [-]
10 cm ² cone	0.14	0.21
15 cm ² cone	0.12	0.17
Ball penetrometer (1)	0.11	0.17
Ball penetrometer (2)	0.13	0.11

NEN-EN-ISO 22476-1: accuracy required for class 2 CPT(U) is 100 kPa (0.1 MPa)

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Permeability

- Unloaded peat has an open structure
- Free water in (large) pores, bounded pore water in cell structure
- When loaded a rapid decay in permeability due to closing of the macro pores



Photo G. Erkens



Permeability

In situ Falling Head test

[m/s]	[m/d]	Depth [m]
8.0×10 ⁻⁷	0.073	1.80
2.3×10 ⁻⁶	0.199	3.49
7.0×10 ⁻⁷	0.060	3.92

From laboratory tests



From literature



Coefficient of Permeability, k_v, m/s



Mesri & Aljouni 2007

Compression indices

17 oedometer tests from test site

	min	max	average
w ₀ [%]	518	1236	875
e ₀ [-]	3.81	8.14	6.19
CR [-] C _c	0.39 2.53	0.58 4.04	0.45 3.21
C_{α}/CR	0.06	0.14	0.08
RR/CR	0.07	0.17	0.13

From literature









CIUC tests



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CIU tests



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CIU tests


DSS, OC, $\sigma_{vc} = 50$ kPa



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DSS, Field stresses



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Anisotropy



Figure 11. Anisotropy of peat revealed by shrinkage tests, Polder Zegveld.

Den Haan & Kruse (2007)

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Linear strain - Natural strain



Oedometer test results in linear and natural strain



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Isotach model based on natural strain, Den Haan (1994) in incremental form



1. Direct strain:

$$\varepsilon_d^H = a \ln\left(\frac{\sigma_i}{\sigma_{i-1}}\right)$$

2. Equivalent time τ directly after loading

$$\tau_{i} = \tau_{i-1} \left(\frac{\sigma'_{i-1}}{\sigma'_{i}} \right)^{\frac{b-a}{c}}$$

3. Equivalent time τ end load step

$$\tau_{i,end} = \tau_i + \Delta t$$

4. Viscous strain

$$\varepsilon_{s}^{H} = c \ln \left(\frac{\tau_{i,eind}}{\tau_{i,begin}} \right)$$

5. Total strain and settlement

$$\varepsilon^{H} = \varepsilon^{H}_{d} + \varepsilon^{H}_{s}, \quad \Delta h = h_{0} \left(1 - \exp(-\varepsilon^{H}) \right)$$
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Results from 107 CRS tests



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Results from 93 CRS tests



Case: Railway line Rotterdam – Germany



Soil profile



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Calculation vs measurements



Vertical 1 (X = -5.317 m; Z = 0.000 m) Method = Isotache with Terzaghi (Natural strain) Fit factors used Coefficient of determination = 0.994 Depth = 1,430 (-) [m Settlement after 10000 days = 2.035 [m]

Long term settlement

Dike with medieval origin



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TO BE SOLVED

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Creep after unloading modeling oedometer test, OCR = 1,8



20 JUPIOINDOL 2017

Creep after unloading modeling oedometer test, OCR = 1,8



Further research

- Creep in 3D, little experience for modeling peat related problems



Influence of gas

Gas is not considered, but plays a role



Den Haan & Kruse 2007



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Creep of Geomaterials



Rate-dependency based EVP modelling approach for clays: from 1D to 3D

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15 September 2014

[PIAP-GA-2011-286397] 2nd CREEP Workshop & School, NTNU

Outline

1. Background-*Phenomena and problems*

2. Rate-dependency of clay

- (1) 1D + remolded clay
- (2) *1D* + *intact clay*
- (3) 3D + remolded clay
- (4) 3D + intact clay
- (5) Applications

3. Conclusions











Clay – Natural hazards





Displacement ___ Stable? Sliding of clay slope Time

Sample scale

Micro scale



Outline

1. Background

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(1) 1D + remolded clay
(2) 1D + intact clay
(3) 3D + remolded clay
(4) 3D + intact clay
(5) Applications

LIM 15.0kV 10.0mm x3.00k SE(M)

3. Conclusions



Strain-rate dependency



Strain-rate dependency

Expon1.
$$\hat{\varepsilon}_{v} = \mu \exp\left[N\left(\frac{p_{m}^{d}}{p_{m}^{s}}-1\right)\right]$$

Expon2.
$$\hat{\varepsilon}_{v} = \mu \left\{ \exp \left[N \left(\frac{p_{m}^{d}}{p_{m}^{s}} - 1 \right) \right] - 1 \right\}$$

$$\mathbf{Power1} \approx \hat{\mathbf{\epsilon}}_{v} = \mu \left(\frac{p_{m}^{d}}{p_{m}^{s}}\right)^{N} \approx$$

$$\mathbf{Power2}_{\varphi} \quad \hat{\varepsilon}_{\nu} = \mu \left[\left(\frac{p_m^d}{p_m^s} \right)^N - 1 \right]_{\varphi}$$

Power3.
$$\hat{\epsilon}_{v} = \mu \left(\frac{p_{m}^{d}}{p_{m}^{s}} - 1\right)^{N}$$





Yin Z-Y*, Karstunen M, Hicher PY. *Soils and Foundations*, 2010, 50(2): 203-214.

Strain-rate dependency



Strain-rate dependency



Yin Z-Y*, Chang CS, Karstunen M, Hicher PY. *International Journal of Solids and Structures*, 2010, 47(5): 665-677.

□ Strain-rate dependency







Simple & easy to determine

□ Rate-dependency ↔ Creep?

Simulations by given values of parameters:



Rate-dependency formulation

Creep of Vermeer



High loading-rate

loading-rat

Stress

relaxatior




Uniqueness of time-dependencies







Yin Z-Y*, Zhu QY, Yin JH, Ni Q. Stress relaxation coefficient and formulation for soft soils. Géotechnique Letters, 2014, 4(1): 45-51.

Outline

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Inter-particle bonds and debonding





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Equations of 1D model



Validation (natural soft clays)



1D-EVP model for natural soft clays. *Yin et al.* (2011)

Outline

1. Background

2. Rate-dependency of clay

(1) 1D + remolded clay
(2) 1D + intact clay
(3) 3D + remolded clay
(4) 3D + intact clay
(5) Applications

3. Conclusions





□ Elastoplasticity vs. Elasto-viscoplasticity

$$d\varepsilon_{ij} = \dot{\varepsilon}_{ij}$$
Strain increments
(EP model)

$$\dot{arepsilon}_{ij} = \dot{arepsilon}_{ij}^{e} + \dot{arepsilon}_{ij}^{p}$$

 $\dot{arepsilon}_{ij}^{p} = d\lambda rac{\partial g}{\partial \sigma_{ij}}$

 $d\varepsilon_{ij} = \dot{\varepsilon}_{ij} \cdot dt$

Strain rates

(EVP model)



 $\log \sigma_{\rm v}$

Elastoplasticity:

- Yield surface (f)
- Flow rule (f=g)
- Hardening rule (κ^*)



Elasto-viscoplasticity:

- Static/reference surface
- Flow rule
- Hardening rule
- Scaling function $(d\lambda)$



(1) Scaling function with flow rule

$$\dot{\varepsilon}_{ij}^{vp} = d\lambda \frac{\partial g}{\partial \sigma_{ij}} = \mu \langle \Phi(F) \rangle \frac{\partial g}{\partial \sigma_{ij}}$$

Equation from 3D to 1D:

Equation in 1D:

$$\dot{\varepsilon}_{ij}^{vp} = d\lambda \frac{\partial g}{\partial \sigma_{ij}} \xrightarrow{underK_{0}} \dot{\varepsilon}_{v}^{vp} = d\lambda \left(\frac{\partial g}{\partial p'}\right)_{K0} \quad \Leftrightarrow \quad \dot{\varepsilon}_{v}^{vp} = \dot{\varepsilon}_{v}^{r} \frac{\lambda - \kappa}{\lambda} \left(\frac{\sigma_{v}'}{\sigma_{p}^{r}}\right)^{\beta}$$
$$d\lambda = \dot{\varepsilon}_{v}^{r} \frac{\lambda - \kappa}{\lambda} \left(\frac{p_{c}^{d}}{p_{c}^{r}}\right)^{\beta} \frac{1}{\left(\frac{\partial g}{\partial p'}\right)_{K0}} \quad \Leftrightarrow \quad \left[\begin{array}{c} \mu = \dot{\varepsilon}_{v}^{r} \frac{\lambda - \kappa}{\lambda} \left(\frac{\partial g}{\partial p'}\right)_{K0}^{-1} \\ \left(\frac{\partial g}{\partial p'}\right)_{K0}^{-1} \end{array}\right]$$
$$\left[\dot{\varphi}(F) = \left(\frac{p_{c}^{d}}{p_{c}^{r}}\right)^{\beta} \frac{\partial g}{\partial \sigma_{ij}} \right]$$



Rate-dependency from 1D to 3D?





Uniqueness of rate-dependency for σ'_{p0} and q_u . Yin et al. (2013)

□ Reference surface?



(2) Reference surface

$$f_r = \frac{q^2}{M^2} + p\left(p - p_c^r\right)$$

(3) Flow rule

$$g = f_{d}$$

$$\frac{\dot{\varepsilon}_{v}^{p}}{\dot{\varepsilon}_{d}^{p}} = \frac{M^{2} - \eta^{2}}{2\eta}$$

(4) Hardening rule

$$\dot{\sigma}_{v} = \sigma_{v} \frac{1 + e_{0}}{\lambda - \kappa} \dot{\varepsilon}_{v}^{p} \Longrightarrow \dot{p}_{c} = p_{c} \frac{1 + e_{0}}{\lambda - \kappa} \dot{\varepsilon}_{v}^{p}$$

Stress history effect on reference surface



□ Mod.-1: Reference surface formulation

$$f_{r} = \frac{q^{2}}{M^{2}} + p(p - p_{c}^{r}) \implies f_{r} = \frac{(q - p'\alpha)^{2}}{M^{2} - \alpha^{2}} + p(p - p_{c}^{r})$$

Initial reference surfaces of some typical natural soft clays



- ✓ Yin et al (2013), Marine Georesources & Geotechnology.
- ✓ Yin et al (2011), ASCE Journal of Geotechnical and Geoenvironmental Engineering.
- ✓ Yin et al (2010), International Journal of Solids and Structures.
- ✓ Karstunen & Yin (2010), Geotechnique.



□ Mod.-2: Reference surface rotation



Initial Anisotropy

α Changing with stresses & strains

Avantage: No additional parameters!

Parameters of ANICREEP model

$$e_0, \kappa, \lambda, \sigma_{p0}, M_c$$
 Parameters of MCC
 $C_{\alpha e}$ or β or R_{α} Viscosity



An anisotropic elastic-viscoplastic model for soft clays

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□ Test simulation (St-Herblain clay)



□ Test simulation (Wenzhou clay, China)





Remarks: Reference surface and rotation?

Pestana & Whittle (1999)



 $\overline{\sigma'_{mb}}$

O.C.Region

 $\sigma'_{mc} \sigma'_m$

Refrence ₽	Surface rotational law.					
Dafalias (1986)₽	$\dot{\alpha} = \left\langle L \right\rangle \frac{1 + e_{in}}{\lambda - \kappa} \left \frac{\partial f}{\partial p} \right c \frac{p}{p_0} (\eta - x\alpha) \phi$					
Whittle (1993)₀	$(\dot{\alpha}\eta_{konc}) = \frac{b}{p_c} \left[\frac{k - \alpha\eta_{konc} - \xi }{k} \right] (q - p'\alpha\eta_{konc}) \dot{\varepsilon}_v^p \leftarrow$					
Hueckel & Tutumluer (1994).	$(\dot{\alpha}\eta_{konc}) = b \left[\frac{q}{p} \exp(-b\varepsilon_s^p) \right] \dot{\varepsilon}_s^p + r\dot{\varepsilon}_r^p _{q^2}$					
Newson & Davies(1996)+	$\dot{\eta}_0 = \pm \left[1 - (\eta_s / M)^2\right] \dot{p} \exp\left(\eta_s - \eta_{os}\right) \zeta \leftrightarrow$					
Hashiguchi (1998) ₄ ,	$\dot{\alpha} = b \left\ \eta - \alpha \right\ \left(\frac{\eta - \alpha}{\left\ \eta - \alpha \right\ } m_b - \alpha \right) \dot{\varepsilon}_{z}^{p}$					
Wheeler et al (2003)¢	$\dot{\alpha} = \mu \left[\omega_d \left(\frac{s_{ij}}{3p'} - \alpha \right) \left \dot{\varepsilon}_d^p \right + \left(\frac{3s_{ij}}{4p'} - \alpha \right) \left\langle \dot{\varepsilon}_v^p \right\rangle \right] $					
Kobayashi et al (2003)¢	$\dot{\alpha} = \frac{J}{D} \left[\frac{(\beta_z + \beta_z \eta - 2 \beta_z \alpha)}{2\ \eta - \alpha\ } \right] \sqrt{\frac{2}{3}} \ D_z^p\ $					
Oka et al (2004)+	$\dot{\alpha} = B_s + (B_0 - B_s) \exp(-B_t \varepsilon_d^p) (A \dot{\varepsilon}_v^p - \alpha \dot{\varepsilon}_d^p) \phi$					
Zhang et al (2007)&	$\dot{\alpha} = \frac{J}{D} b_r (b_l M - \alpha) \sqrt{\frac{2}{3}} \ D_z^p \ \frac{\hat{\eta}}{\ \hat{\eta}\ }$					
ч	²					

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Outline

1. Background

2. Rate-dependency of clay

(1) 1D + remolded clay
(2) 1D + intact clay
(3) 3D + remolded clay
(4) 3D + intact clay
(5) Applications



3. Conclusions

□ Extension of debonding rule from 1D to 3D



Time-dependent modeling of soft sensitive clay, Yin et al (2011)

□ Parameters of ANICREEPS model



ASCE AMERICAN SOCIETY JOURNAL OF GEOTECHNICAL AND GEOENVIRONMENTAL ENGINEERING © ASCE / NOVEMBER 2011 / 1103 Modeling Time-Dependent Behavior of Soft Sensitive Clay

Zhen-Yu Yin¹; Minna Karstunen²; Ching S. Chang, M.ASCE³; Mirva Koskinen⁴; and Matti Lojander⁵ 42

Test simulation (Vanttilla clay, Finland)

$\tau = 1 \text{ day (Conventional)}$

$\tau = 10 \text{ days}$

$\tau = 100 \text{ days}$



Each load increment lasts 1 d、 10 d、 100 d!

Undrained triaxial CRS tests

Test simulation (Vanttilla clay, Finland)





Undrained triaxial creep tests

Outline

1. Background

2. Rate-dependency of clay

- (1) 1D + remolded clay
 (2) 1D + intact clay
 (3) 3D + remolded clay
 (4) 3D + intact clay
- (5) Applications

3. Conclusions





□ Implementation of model into PLAXIS (user defined model)

🎬 Plaxis 8.2 In	nput - tunnel.plx		-									
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General Parameters Interfaces												
	Material Set General Parameters Interfaces											
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□ Haarajoki embankment, Finland





Haarajoki test embankment

All parameters determined from conventional oedometer and triaxial tests:



□ Finite element model



Embankment fill (Mohr-Coulomb)

Thickness (m)	$E (kN/m^2)$	υ'	<i>ø</i> '	ψ'	$c'(kN/m^2)$	γ (kN/m ³)
0-3	40 000	0.3	38 °	0 °	1	21

Excess pore pressure



Total displacement



IkN/m²I

6.000

34 000

-40.000 -42.000 -44.000







Construction time effect on excavation

Total displacement



Pile construction (installation rate effect, by Abaqus)

Excess pore pressure



Effect of explosion on clay excavation (terrorist attack, by LS-Dyna)

Contour (Analysis system) Time = 0.000000 Displacement (Mag) -7.135E-03 6.342E-03 -5.549E-03 -4.757E-03 -3.964E-03 3.171E-03 2.378E-03 1.586E-03 7.928E-04 0.000E+00 No result Max = 7.135E-03 Min = 0.000E+00

Total displacement
3 Conclusions

- (1) Determination of 1D strain rate-dependency formulation;
- (2) From 1D to 3D, attention to flow rule, reference surface, anisotropy, destruction;
- (3) Applications on geotechnical structures.
- **Main collaborators**
 - ✓ Pierre-Yves HICHER (GeM-ECN)
 - Minna KARSTUNEN (UT-Chalmers, Suède)
 - Ching S. CHANG (Umass, USA)
 - ✓ Jian-Hua YIN (HK-PolyTechU)



Related publications

- Yin Z-Y*, Xu Q, Yu C. Elastic viscoplastic modeling for natural soft clays considering nonlinear creep. *ASCE International Journal of Geomechanics*, <u>doi: 10.1061/(ASCE)GM.1943-5622.0000284</u>.
- Wang L-Z, Yin Z-Y*. Stress-dilatancy of natural soft clay under undrained creep condition. *ASCE International* Journal of Geomechanics, doi: 10.1061/(ASCE)GM.1943-5622.0000271.
- Yin Z-Y*, Yin JH, Huang HW. Rate-dependent and long-term yield stress and strength of soft Wenzhou marine clay: experiments and modeling. *Marine Georesources & Geotechnology*, DOI:10.1080/ 1064119X.2013.797060.
- Yin Z-Y*, Zhu QY, Yin JH, Ni Q. Stress relaxation coefficient and formulation for soft soils. *Géotechnique Letters*, 2014, 4(1): 45-51.
- **Yin Z-Y***, Karstunen M, Chang CS, Koskinen M, Lojander M. Modeling time-dependent behavior of soft sensitive clay. *ASCE Journal of Geotechnical and Geoenvironmental Engineering*, 2011, 137(11): 1103-1113.
- **Yin Z-Y***, Karstunen M, Hicher PY. Evaluation of the influence of elasto-viscoplastic scaling functions on modelling time-dependent behaviour of natural clays. *Soils and Foundations*, 2010, 50(2): 203-214.
- Yin Z-Y*, Chang CS, Karstunen M, Hicher PY. An anisotropic elastic viscoplastic model for soft clays. International Journal of Solids and Structures, 2010, 47(5): 665-677.
- Karstunen M, Yin Z-Y*. Modelling time-dependent behaviour of Murro test embankment. *Géotechnique*, 2010, 60(10): 735-749.
- Yin Z-Y*, Hicher PY. Identifying parameters controlling soil delayed behaviour from laboratory and in situ pressuremeter testing. *International Journal for Numerical and Analytical Methods in Geomechanics*, 2008, 32(12): 1515-1535.



Thanks a lot for your attention

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Use of advanced creep models and some pitfalls in creep modelling

G. Grimstad

Fundamental aspects of soft clay behavior

- Creep
- Anisotropy
 - Strength
 - Stiffness
 - Yield stress
- Structure and destructuration
- Unloading/reloading cycles small strain
- Degradation during cyclic loading
- ALL ARE LINKED!

"1D" Creep – (24h) incremental oedometer test

- Advantages:
 - Gives first estimate of creep/consolidation parameters and the "vertical" pre-consolidation stress directly
- Disadvantages
 - Time consuming compared to CRS tests
 - Only average settlement parameters for large stress increments
 - Ideally back calculation with mathematical model is needed (FEA)

Current Norwegian engineering practice

- Using low OCR (if material has not been subjected to preloading an OCR of 1.0? is often used)
- Ignoring creep
- Adding creep after consolidation?(Hyp. A!)
- Advanced: Janbu's time resistance concept
- What about the selected pre-consolidation stress? IMPORTANT!!!
- What about sample quality?

Sample quality



Compression curves for Väsby clay at a depth of 4.0 – 4.3 m (after Leroueil and Kabbaj, 1987)



DeGroot et. al. (2005)

Janbu's time resistance concept Time_ Increment in time lacksquaredivided by the increment in strain (Cause/Effect). ε_a $R = \partial t / \partial \varepsilon$ "Pure creep" r_s **R**_{ref} X Time

τ

1D equation



"Alternative approaches" to Janbu for 1D

- Yin and Graham (equivalent time approach) Adopted from Bjerrum
- Leroueil
- Den Haan (ABC model)
- etc.

ALL ARE THE SAME?



A case of SSC and SS model giving the same final settlement.



Illustration of dependence of OCR on the corresponding reference time (τ).

The effect of the $\mu^*/(\lambda^*-\kappa^*)$ ratio on OCR (creep rate)

			_
$\frac{\mu^*}{\lambda^* - \kappa^*}$	age = 10000 yrs	OCR = 1.3	
$=\frac{1}{r_s\cdot\zeta}$	OCR	age	
0.010	1.163	6.79E+08 years	
0.015	1.254	1.08E+05 years	
0.020	1.353	1.36E+03 years	
0.025	1.459	98.9 years	
0.030	1.574	17.2 years	
0.035	1.697	4.93 years	$\left \frac{1}{H} \right $
0.040	1.830	1.93 years	
0.050	2.129	0.518 years	
0.070	2.880	0.113 years	



"Recommended" range (PLAXIS manual) ~0.04 - 0.07

Anisotropy



- First:
 - Undrained Triaxial Compression versus Undrained Triaxial Extension and Direct Simple Shear (Bjerrum 1973)
- Second:
 - Preconsolidation stress from Oedometer test versus isotropic consolidation test (Feng 1991)
- Third:
 - "Stress/strain induced anisotropy" Changes in macroscopic yield surface (Wheeler 2003)

Undrained shear strength

 Used as basis for the NGI-ADP model



Pre-consolidation stress and "cap" yield surface

 Experiments from literature on finding cap surface – yield points in p' – q space



Stress/strain induced anisotropi 30 30 CAE2513 CAE2544 Wheeler η=-0,66 η=-0.59 \bullet α=-0.15 α=-0.11 et al. 10 10 q (kPa) q (kPa) 20 40 20 p′ (kPa) p' (kPa) -10 -10 -30 -30 30 40 CAE2496 CAD2276 η=-0.35 η=0.25 α=-0.11 α=0.15 20 10 q (kPa) q (kPa) 0 20 40 20 40 60 80 p' (kPa) -10 p' (kPa) -20

-40

-30

Destructuration



Burland (1990)

Creep - Yield surface becomes reference surface

- Option 1 extending by volume strain (ACM)
- Option 2 extending by plastic multiplier directly

$$\dot{\lambda} = \frac{1}{r_s \cdot \tau} \cdot \left(\frac{p^{eq}}{p_{ref}^{eq}}\right)^{r_s \cdot \zeta} \cdot m_{KONC}$$



Anisotropy and creep – The n-SAC model

- A non-associated creep model for structured anisotropic clay
- Non-associated because:
 - prediction of the strain behavior under various stress paths, based on experimental evidence from e.g. Feng (1991)

$$p^{eq} = p' + \frac{\frac{3}{2} \{\mathbf{\sigma}_d - p' \mathbf{\beta}_d\}^T \{\mathbf{\sigma}_d - p' \mathbf{\beta}_d\}}{\left(M^2 - \frac{3}{2} \mathbf{\beta}_d^T \mathbf{\beta}_d\right) p'}$$
$$Q = p' + \frac{\frac{3}{2} \{\mathbf{\sigma}_d - p' \mathbf{\alpha}_d\}^T \{\mathbf{\sigma}_d - p' \mathbf{\alpha}_d\}}{\left(M_f^2 - \frac{3}{2} \mathbf{\alpha}_d^T \mathbf{\alpha}_d\right) p'} - p_Q^{eq} = 0$$

where p' = mean stress; σ_d =deviatoric stress vector; β_d = deviatoric rotational vector; M = Lode angle dependent peak of the reference curve of in p'-q space

where M_f is the Lode angle dependent citical state line in p'-q space; $\mathbf{\alpha}_d$ is the deviatoric rotational vector.

Modelling of destructuration



- Gens and Nova (1993)
- *p_{ref} = p_{mi}* ·(1+x)

 where x is the amount of structure that is unstable

 i stands for intrinsic

$$\dot{\lambda} = \frac{1}{r_{si} \cdot \tau} \cdot \left(\frac{p^{eq}}{(1+x) \cdot p_{mi}'}\right)^{r_{si} \cdot \zeta_i} \cdot m_{K0N0}$$

• x has to change with vp strain $\frac{dx}{d\lambda} = f(\Phi)^{-\text{State variables inc. }x}$

Models with anisotropy and destructuration



- Option 1
 - ACM -> ACM-S (Leoni 2008, Kamrat-Pietraszewska 2011)
 - Extension of SSC (Stolle et al. 1999) (PLAXIS current model)
- Option 2
 - EVP-SCLAY1S (Karstunen and Yin 2010)
 - Ani-Creep (Yin et al. 2011)
 - n-SAC (Grimstad et al. 2010)
- n-SAC –using creep limit and option 2:

$$\dot{\lambda} = \frac{1}{r_{si} \cdot \tau} \cdot \left\langle \left(\frac{p^{eq}}{(1+x) \cdot p_{mi}} \right)^{r_{si} \cdot \zeta_i} - \frac{\tau}{t_{max}} \right\rangle \cdot m_{K0NC}$$

How to use/Parameters for analyses

- Two models SSC and n-SAC
- Three analysis cases SSC1, SSC2 and n-SAC

Model	ν	K ₀ ^{NC}	E _{ref} / p _{ref}	${E_{oed}}^{ref}_i / p_{ref}$	r _{smin}	r _{si}	ω	φ _p	φ _{cs}
SSC1	0.15	0.54	200	9.5	-	267	-	-	35°
SSC2	0.15	0.54	200	6.0	-	233	-	-	35°
n-SAC	0.15	0.5	200	13.0	200	625	0.3	25°	35°

 $k_v = k_h = 5e-5 \text{ m/day}; \gamma' = 10 \text{ kN/m}^3, K_0 = 0.54, \text{OCR} = 1.36$

$$\kappa = \frac{3(1-2\nu)}{E_{ref}} \qquad \qquad \lambda = \frac{1}{\left\{E_{oed}^{ref}\right\}_i} \qquad \qquad \mu^* = \frac{1}{r_s}$$



Example: settlement problem







0.80



Mesh dependency due to softening

Shadings of "structure"





25.7 days. Looks more like perfectly plastic behavior!

Conclusions

- Creep/rate and anisotropy are important if we want to fully understand soil behavior.
- Sample quality is crucial and deserves more attention as it forms the basis for numerical modeling.
- With increased sample quality and testing procedure, the soil models also needs to be improved
- The "huge gap" between state-of-the-art and stateof-the-practice must be closed or at least narrowed down!

Use of creep models

- Expected new stress state to cross p_c' (24h)
- Expected new stress state below p_c' (24h)



Time resistance concept



- Described by in e.g. Janbu (1969)
- Used for 1D strain in KRYKON, Svanø & Emdal (1986)



Example: SSC model - The effect of the $\mu^*/(\lambda^*-\kappa^*)$ ratio on OCR (creep rate)



The effect of the $\mu^*/(\lambda^*-\kappa^*)$ ratio on OCR (creep rate)

			$\frac{H}{H_0}$ λ^* decreasing with stress
$\frac{\mu^*}{\lambda^* - \kappa^*} = \frac{1}{r_s \cdot \zeta}$	$OCR_{\tau} = \left(\frac{t}{\tau}\right)^{\frac{\mu^*}{\lambda^* - \kappa^*}}$ $\tau = 1 \text{ day}$ $t = 10^4 \text{ years}$	$t_{age} = \tau \cdot OCR_{\tau}^{\frac{\lambda^* - \kappa^*}{\mu^*}}$ $\tau = 1 \text{ day}$ $OCR_{\tau} = 1.3$	IO^{3} Ve_{ars} $In(\sigma')$ Same u*
0.010	1.163	6.79E+08 years	$\frac{H}{H} \uparrow \sigma_{v0}' \sigma_{vc 24h}$
0.015	1.254	1.08E+05 years	
0.020	1.353	1.36E+03 years	
0.025	1.459	98.9 years	Vears
0.030	1.574	17.2 years	$ In(\sigma') $
0.035	1.697	4.93 years	$\frac{H}{H_0}$ Same Λ^*
0.040	1.830	1.93 years	
0.050	2.129	0.520 years	LUT KEQIS
0.070	2.880	0.116 years	μ^* decreasing with time
			$\ln(\sigma')$

"Recommended" range (PLAXIS manual) ~0.04 - 0.07

The oedometer test...



- Can we rely on OCR from IL oedometer tests?
 - 1) Sample disturbance? (we all know this)
 - 2) Stress condition?
 - Start at some unknown isotropic stress condition and consolidates to 12.5 kPa of vertical stress
 - Loads further along a line different from K₀^{NC} line (i.e. stress path hits the reference pre-consolidation at different place than it would in-situ!)
 - \rightarrow Do we need to simulate the oedometer test rather then interpret OCR from it?
 - \rightarrow Should we measure horizontal stress in the oedometer?
 - Consolidation (is the effective stress constant for most of the 24h?, e.g. clays with low permeability)
 - 4) Extrapolation... (should model OCR and reality OCR be the same?)

Simulated oedometer with SSC



What about K₀?

- Is the in-situ K₀ affected by creep (NC clay)?
 - Model says: very limited influence, i.e. $K_0 \approx K_0^{NC}$
- Has the material been unloaded (OC clay)?
 - Model says: yes, but creep will try to make K₀ ≈ K₀^{NC} if the model is not changing its plastic potential, since the volumetric strain should be equal to the vertical strain
- Should we then set $K_0 \approx K_0^{NC}$ for models like SSC?
OCR and K_0

- The K_0 value does not change significantly in a 1D creep case due to the increase in OCR. Since 1D creep requires $d\epsilon_1^{vp} = d\epsilon_v^{vp}$, then the stress state is fixed to one point at the potential surface.
- In PLAXIS if one specify a OCR (due to creep alone), the suggested initial horizontal stress generated (suggested K₀) is based on the assumption of unloading. Remember to change this back to a value close to the real K₀^{NC}



Stress increment in the field

- No need to fit the whole lab curve...
 - What is the experienced stress change?
 - For most of the soil it is little change (around p_c' or less)



Accept wrong OCR – Fit at large stress change, well above p_c'

Accept that Inital creep rate is too big

In most cases: Fit for the actual stress change in the region around p_c ', higher λ^* gives lower OCR for same μ^*

The MIT–MDPW embankment

 Latest paper looking at back calculating this is from 2012 (Fatahi et al.)





The trial embankment



Alternative models





Oedometer simulations







Results





Time [day]







Conclusions

- Evaluate the parameters over relevant stress increments
- Do not blindly take OCR from odeometer tests
 - Stress path
 - Sample disturbance
- The "simple" SSC model performs OK when we are after vertical deformation profile and pore pressure. As long as we take some care for the OCR we use in modelling.
 - NC clay does not usually have OCR of 1.1...
 - OCR in SSC is a material parameter that defines initial state of the soil (i.e. the state variable p₀^{eq}), it is not more holy than the other parameters that we use to fit our model to "reality"

$$\lambda^*$$
, κ^* , μ^* , ν , φ , c , K_0^{NC}

Creep of Soft Clay: Exercise

Hans Petter Jostad

Discipline leader in numerical modeling at NGI Adjunct Professor at NTNU

2st CREEP Course, Trondheim, Norway,

15-16. September 2014





CREBS

- 4 Workshops on <u>CRE</u>ep <u>Behaviour of Soft clay</u>)
 - NGI (Oslo, Norway, January 2006)
 - Univ. Stuttgart (Pisa, Italy, September 2007)
 - Univ. Chalmers (Gothenburg, Sweden, July 2009)
 - Deltares (Delft, Netherland, January 2014)

Establish a common basis of understanding "long term compaction in soft soil"

- analyse a set of well defined hypothetical cases

Example calculations

- Comparison of results obtained by different calculation programs (for a set of well defined cases)
- Comparison of material models
- Interpretation of laboratory tests (model dependent)
- Recommendations of laboratory tests and field investigation
- Not a competition!

Hypothetical cases

- 1. NC-behaviour (OCR=1)
- 2. NC-behaviour with apparent pre-consolidation
- 3. Varying time history (pre-loaded several years)
- 4. Layered soil profile (different permeability)
- 5. Stress distribution with depth (some shear strain)

The real case: Oslo Railroad Customs Building

- 50 years with measurements (may include additionally 30 years)

6 Participants

- University of Stuttgart
 - Dr. Martino Leoni and Professor Pieter Vermeer
- University of Strathclyde (and Ecole Central de Nantes)
 - Dr. Zhen-Yu Yin and Professor Minna Karstunen
- University of BRISTOL
 - Dr. David Nash
- Chalmers University of Technology (Gothenburg)
 - Mats Olsson and Professor Claes Alén
- Swedish Geotechnical Institute (SGI)
 - Per-Evert Bengtsson and Rolf Larsson
- Norwegian Geotechnical Institute (NGI)
 - Professor II Hans Petter Jostad

Models

- Plaxis (FE) with Soft Soil Creep and Anisotropic Creep (3D)
 - EVP, MCC, rotated modified CC, no structure, one creep parameter
- Plaxis (FE) with EVP SCLAY-1S (3D)
 - EVP, rotated modified CC, over-stress formulation, structure, two creep parameters
- Briscon (FD) with a general isotache model (1D)
 - EVP, structure, stress dependent creep parameter
- Embankco (FD) with an isotache model (1D)
 - EPVP, structure and threshold value for creep
- GeoSuite (FE) with two slightly different isotache models (1D)
 - EPVP, structure by stress dependent creep

Hypothetical case 2

q = 50 kPa (light) and 90 kPa (heavy)
drainage

$$M = 10 MPa \quad \gamma' = 10 \text{ kN/m}^3$$

 $\gamma' = 10 \text{ kN/m}^3$
 $OCR = 1.4 (10 000 \text{ years old})$
 $e_0 = 1.17 (p_0' = 143 \text{ kPa})$
 $k_v = 0.02 \text{ m/year}$
 $Ip = 18-25\%$
Closed bottom

Main assumption

- 1. Fully saturated clay (incompressible pore water)
- 2. 1D Condition
 - a. 1D pore water flow with defined drainage conditions
 - b. Negligible horizontal strains (oedometer condition)
- 3. Uniform material (only changes in stress levels and initial void ratio) within the soil layers
- 4. Assumed "perfect" oedometer test data?

Effect of sample disturbance?

Soft Soil Creep – input parameters



$$R = R_o e^{\left(r \,\Delta \varepsilon_{creep}\right)} = R_o OCR^{\frac{\lambda^* - \kappa^*}{\mu^*}} = \frac{t_{ref}}{\mu^*} OCR^{\frac{\lambda^* - \kappa^*}{\mu^*}}$$



Oedometer results – standard IL test

20 mm sample with drainage at top and bottom



Oedometer test data



Hypothetical cases





Interpretation – reference strain (24 hours)



Odometer test 693



Interpretation - creep phase (NC-regime)





Interpretation - creep phase (NC-regime)





Interpretation - creep phase (OC-regime)





NB! Creep from one increment affects the creep in the next increment



EVP-SCLAY1S - interpretation



Comparison of stress-strain-time curves







Generation of pore pressure due to creep?

 $\frac{\partial u}{\partial t} < 0?$





Results (strain after 50 yr)





Light building

Heavy building

Oslo Railroad Customs Building





Measured Results


Conclusions

- Large differences in settlements for well defined idealized examples
- The main reason is uncertainties in the creep behaviour before the yield stress (apparent pre-consolidation pressure)
- The differences may have been even larger due to uncertainties in the pre-consolidation pressure (if not given!)
- The programs and material models seems to work well (except EMBANCKO?)
- Difficult to check the results obtained with the general 3D models (especially EVP-SCLAY1)
- Difficult to compare models due to different sets of input parameter even when they are based on the same framework



Creep Modelling of Soft Soils

Final remarks

Jelke Dijkstra



- Should we start incorporating kinematic hardening models to capture unloading/reloading + creep (DMW)
 - Already a model in development (Chalmers/NGI)

by Nallatamby Sivasithamparam





 Should we start incorporating kinematic hardening models to capture unloading/reloading + creep (DMW)

Already a model in development (Siva@NGI)

- Fibre overlay model + clay model = peat? (DMW)
- Creep in sand is not only governed by grain crushing (DMW)
- Hypothesis B is the only physical realistic explanation (HPJ)
- Extrapolation of laboratory time to in situ time scales remains challenging: 24 hrs << 50 years (HPJ)

- Long-term field data is not easy to come by either

• Separation of consolidation & creep stage is challenging (HPJ)



- Consider sample disturbance in the determination of the relevant parameters from lab data (HPJ)
 - can we correct oedometer data for apparent preconsolidation pressure?
 - It is disappointing that you only know AFTER the test if your sample is of insufficient quality
- Peat classification is ambiguous (CZ)
 - So let's use the simplified Russian system ...



- Classic approach for calculation of consolidation and creep is surprisingly effective in 1D (CZ)
 - What is creep in peat anyway?
 - Is decay of peat also exponential, such as
 - chemical reactions, electric charge, discharge of fluid from a vessel, atmoshperic pressure, heat transfer, luminescence, biological half lives, electromagnetic radiation, radioactivity, thermoelectricity, damped mechanical oscillators etc.
 - Field data for calibration of predicitons
 - Extension to 2D & 3D required



- Measurement at very low (effective) stress levels are not trivial to perform (CZ)
- Viscoplastic models capture the phenomenological system response very well (Yin)
 - Though not all input parameters are intuitive (or shown)
 - Sensitive for the structure parameters
- Smart use of simplified models will, in some cases, approach advanced models. Determine your parameters around your design point and tell Plaxis the 'age' of the soil (GG)
- Ranking of the most important parameters for Soft Soil Creep OCR, OCR, OCR, and OCR (GG)



Outlook

Should we consider mesoscale modelling?





Matsushima (2014)



Outlook

- Or even on the molecular scale?
- Ebrahimi et al. 2014

