



Creep of Geomaterials

No.286397







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Artificial Ground Freezing Widely used in Europe

Vienna: For the extension of the underground line U2 from the first to the second district of Vienna, the underground station Schottenring is constructed with low overburden right beneath the Danube channel. Because of the low overburden, artificial ground freezing was employed as temporary support and sealing device. (http://www.imws.tuwien.ac.at)

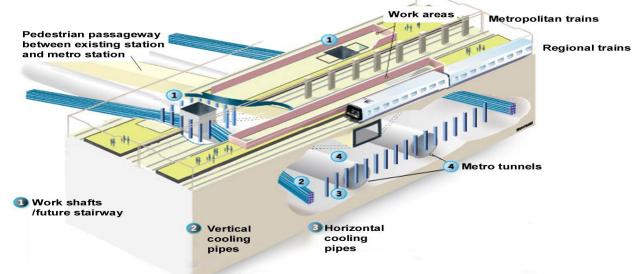
Copenhagen: A pedestrian passage from a new metro station to an existing railway station

was constructed underground. Since the existing rail traffic had to continue, the ground was

frozen to avoid the risk of collapse due to excavation of the transfer tunnel.

Berlin: unter den Linden

Metro Station (2006)



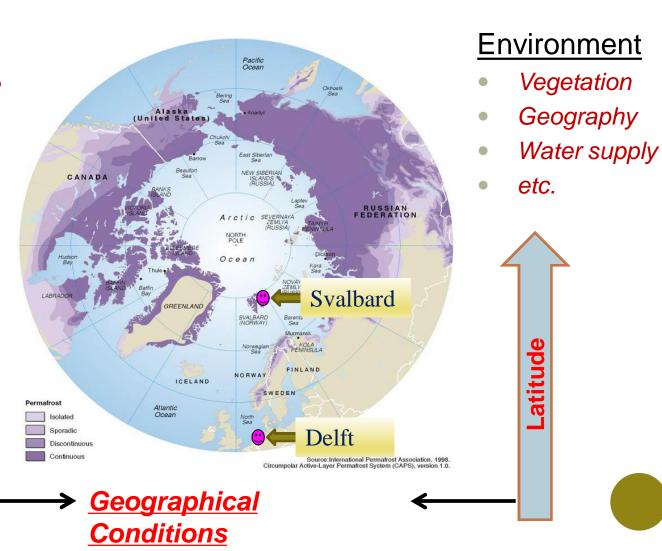
The Netherlands: Artificial Ground Freezing at Sophiaspoortunnel

Soft clay: supporting during excavation; diaphragm wall

Permafrost distribution

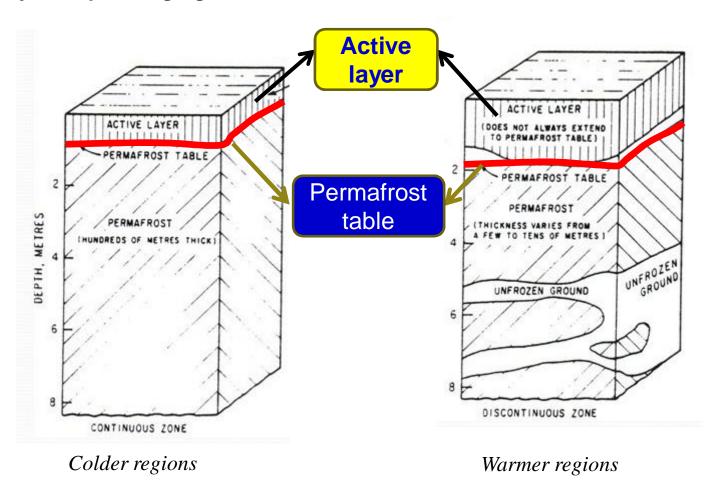
Climate

- Temperature
- Precipitation
- Wind speed
- o etc.



Permafrost table and Active layer

Active layer: depth ranging from tens cm to several metres



From Andersland and Ladanyi (1999)

AGENDA

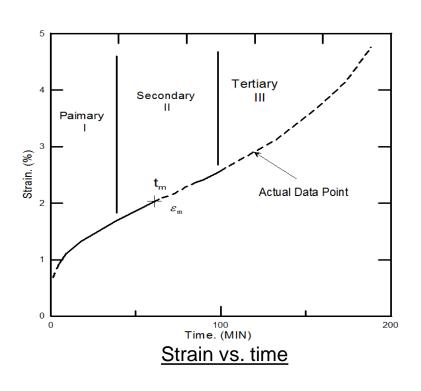
- Factors influencing creep of frozen soils
- □ Creep of Warm Frozen Soils: Field Observations
- State of The Art: Creep Models for Frozen Soils
- Our attempts on constitutive modeling
- Concluding Remarks

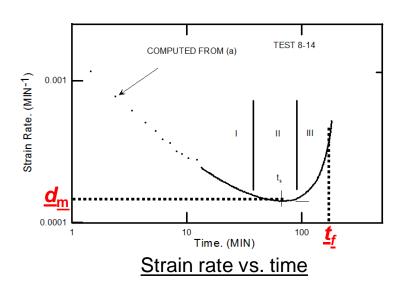
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CREEP CURVES FOR FROZEN SOILS

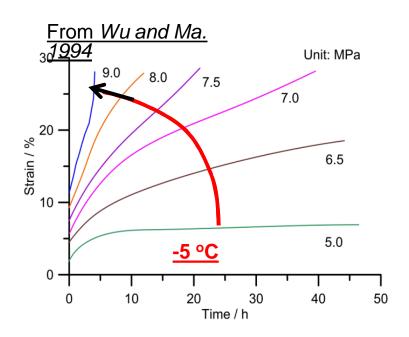
TYPICAL creep curves for frozen soil (Ting 1983)



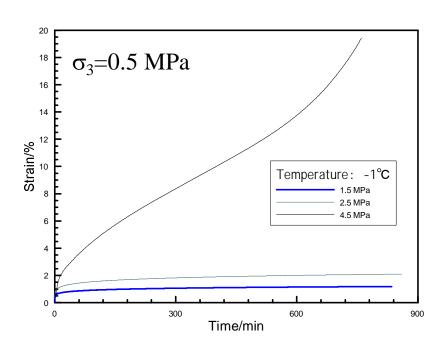


Frozen soil displays features very similar to that of unfrozen soil

STRESS DEPENDENCE



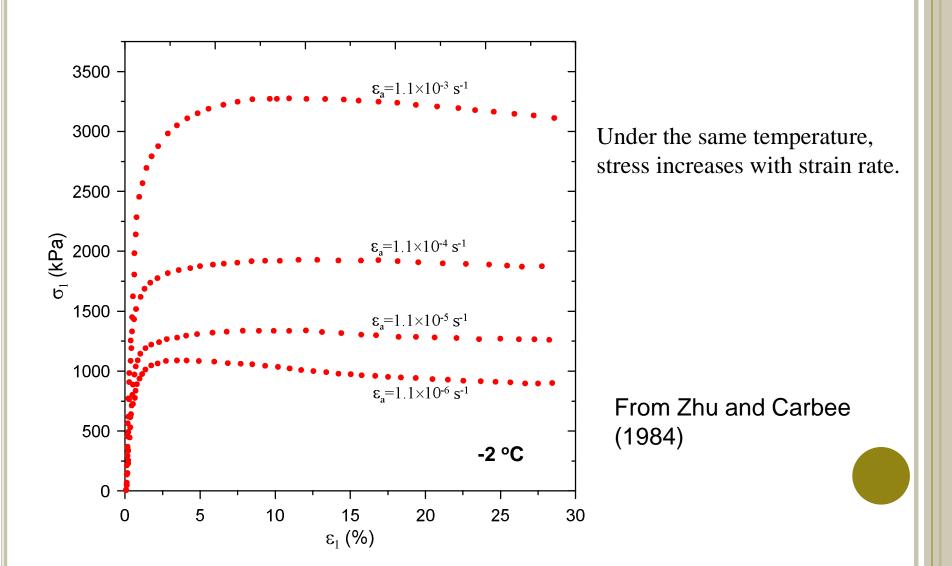
Uniaxial compression Frozen Lanzhou sand



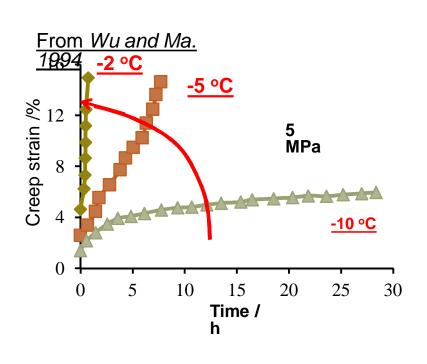
Triaxial compression Frozen ISO Standard sand

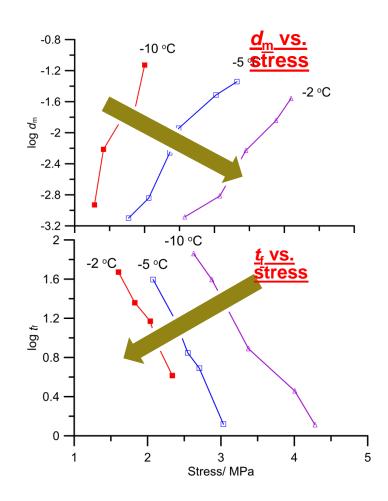
Both very much stress dependent.

STRAIN RATE DEPENDENCE



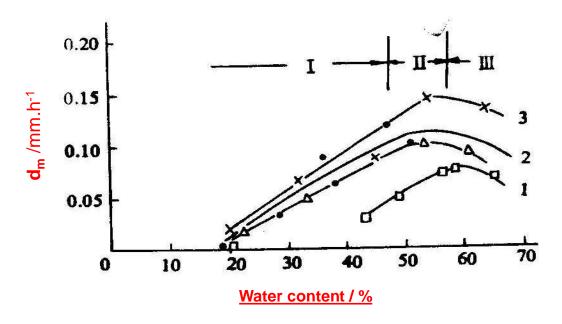
THERMAL DEPENDENCE





Temperature plays a role similar to load Minimal strain rate and time to failure all dependent on temperature

DEPENDENT ON TOTAL WATER CONTENT*.



1. -0.5 °C, 1.4 MPa; 2. -1.0 °C, 2.2 MPa; 3. -2.0 °C, 3.4 MPa

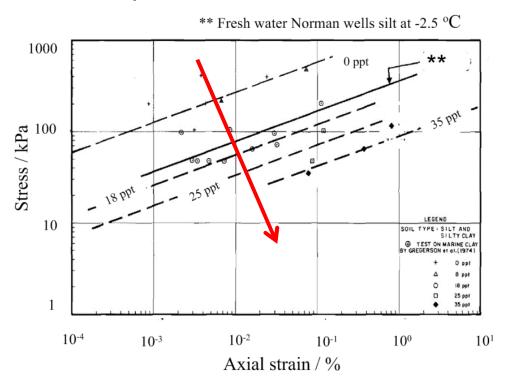
From Wu and Ma. 1994

Generally, different water content range, different minimal strain rate development for frozen silty clay (samples from Lanzhou, China)

* Total water content: ice is counted as water together with unfrozen water

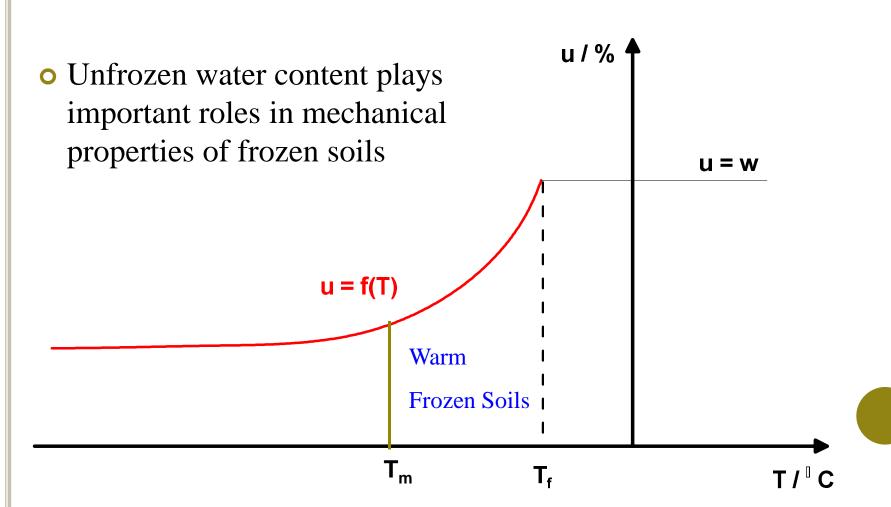
SOLUTE DEPENDENCE

Solute (Nixon and Lem 1984):



The higher solute content, the larger axial strain under a certain stress level.

TEMPERATURE & SOLUTE CHANGE UNFROZEN WATER CONTENT



AGENDA

- Factors influencing creep of frozen soils
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- Concluding Remarks

DEFINE WARM FROZEN SOIL (Qi and Zhang, 2008)

- Shields et al. (1985): -2.5 and -3.0 °C.
- Foster et al. (1991): -1 °C
- Tsytovich(1975): different temperature boundaries for different soils

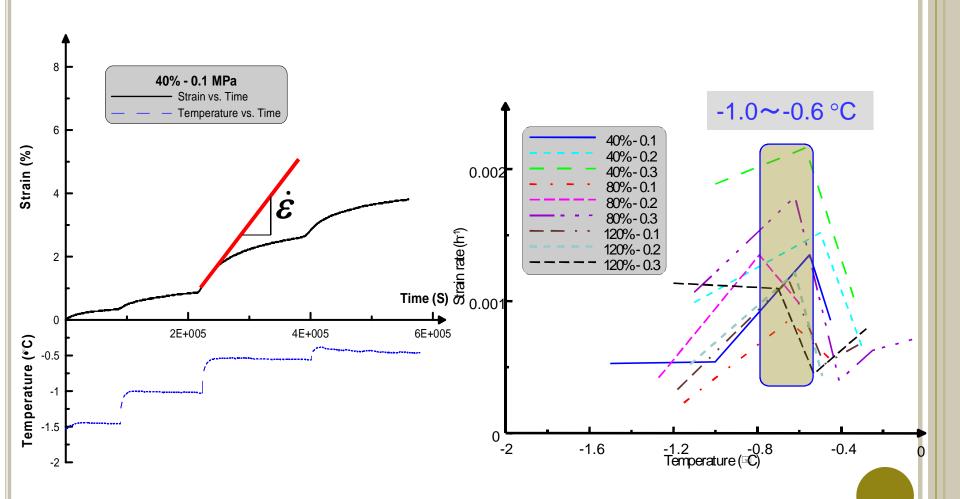
Temperature boundary:

- ➤ Detecting unfrozen water content is rather difficult (NMR)
- > Tend to use mechanical properties

Constant load and stepped temperature

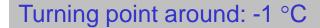
- Material: silty clay (CL) Qinghai-Tibet Plateau
- Different water content: 40%, 80%, 120%
- Constant Loads: 0.1, 0.2 and 0.3 MPa, respectively
- Temperature steps: -1.5, -1.0, -0.6, -0.5 and -0.3 °C

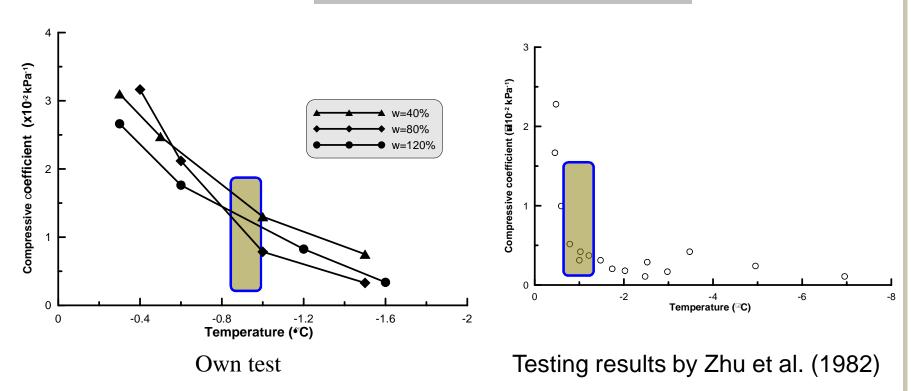
➤ Strain rate vs. Temperature



like we applied higher loads in oedometer test

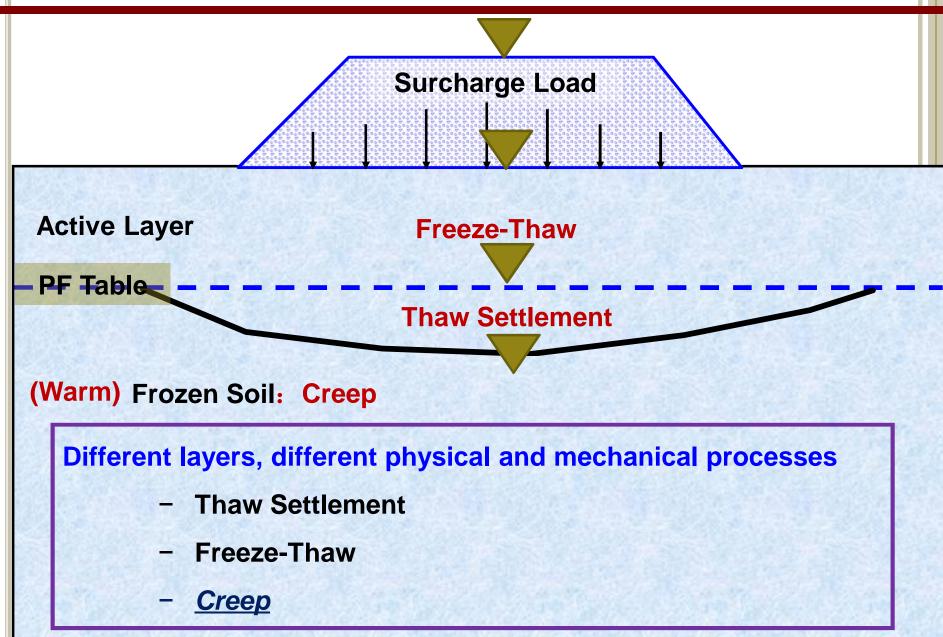
Compressive Coefficient vs. Temperature



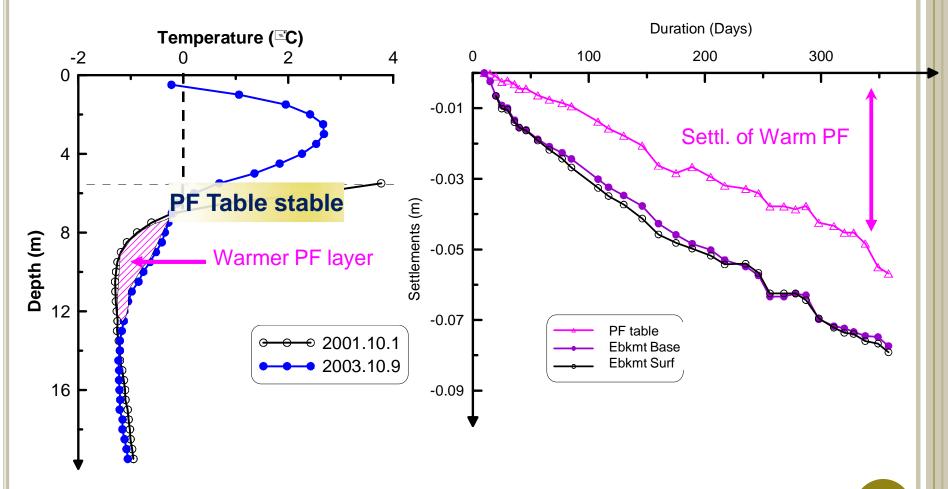


-1 °C is defined as the temperature boundary for warm frozen soil for the silty clay we frequently encounter on the plateau.

Settlement of road embankment

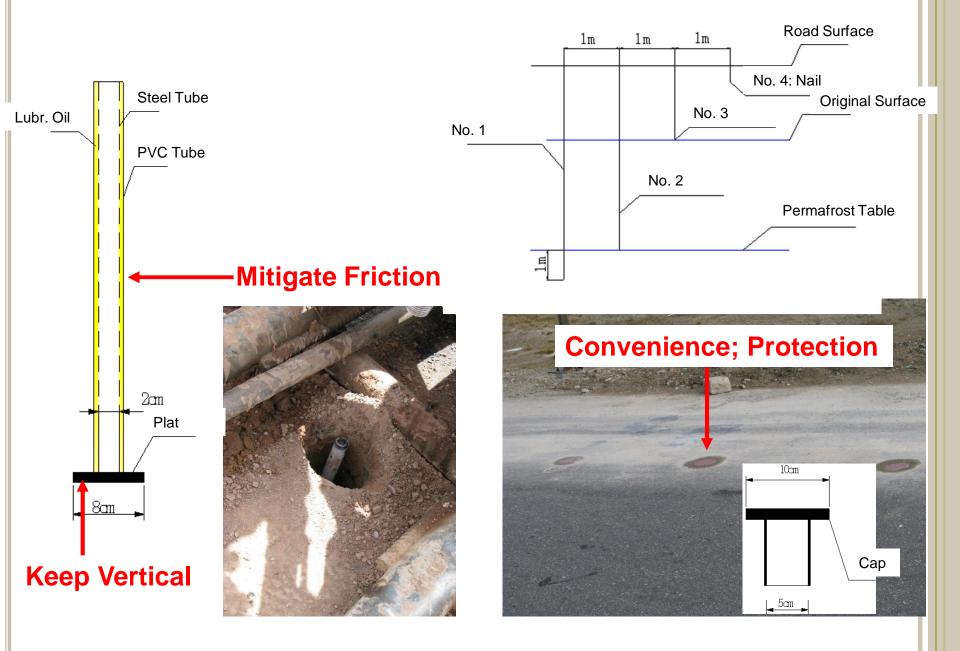


FIELD OBSERVATION QINGHAL-TIBET RAILWAY.

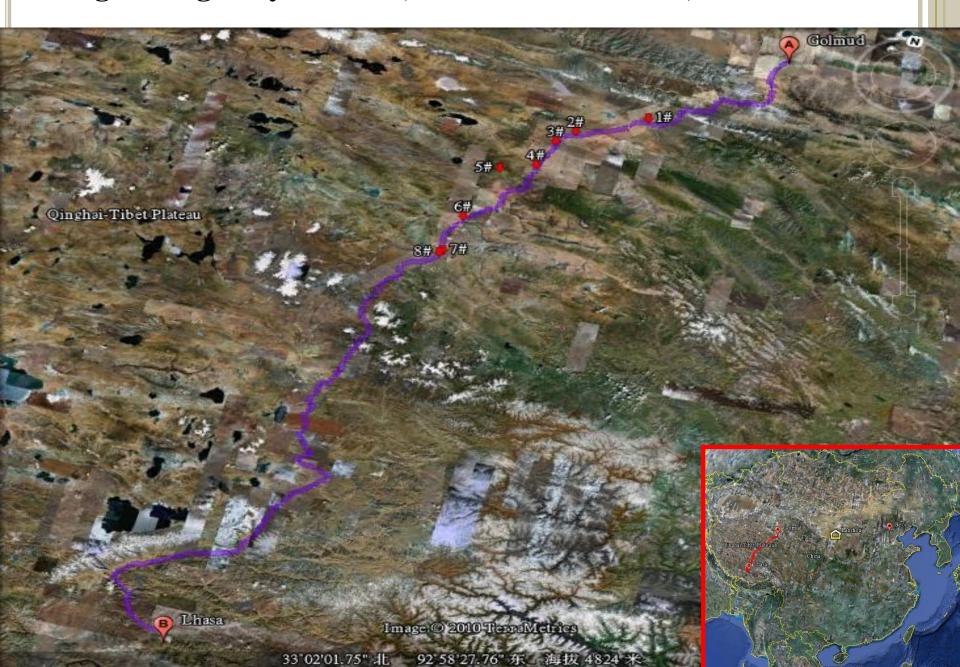


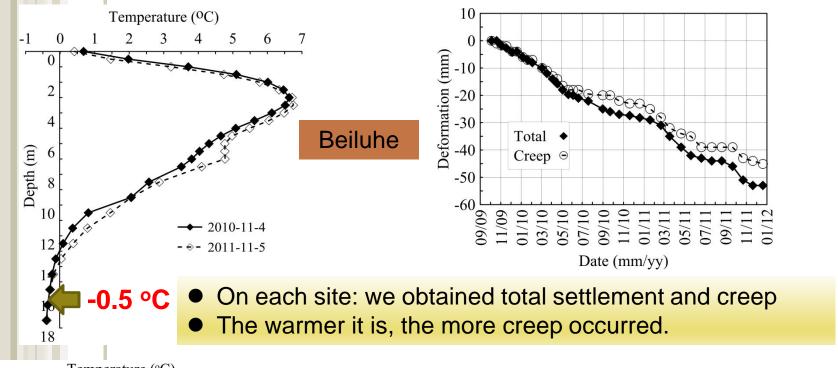
Location: DK 1136+540 (QT Railway)

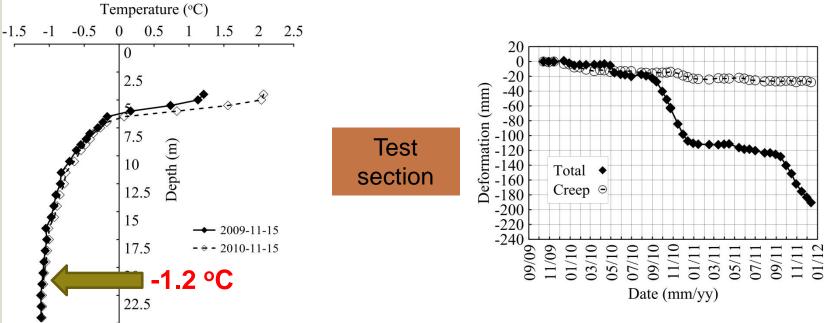
FIELD OBSERVATION QINGHALTIBET HIGHWAY



Along the highway: 300 km, 10 observation sites, 4500 m a.s.l.

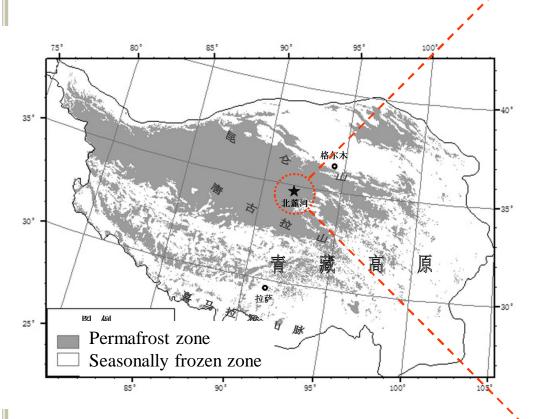






LONG-TERM LOAD TEST.

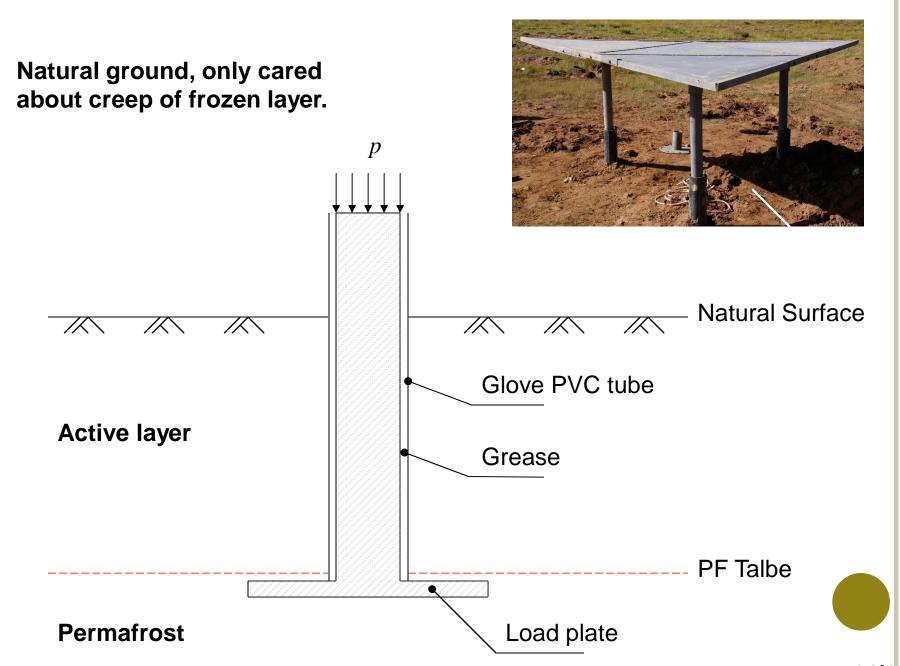
(by Dr. Zhang)

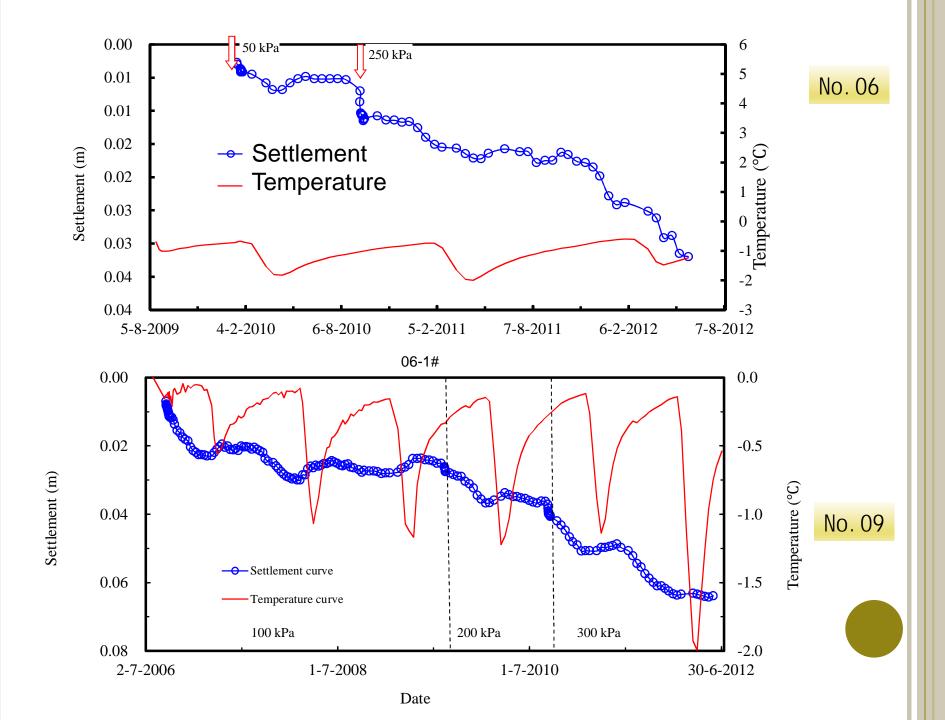


Working Site: Beiluhe field station









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CREEP MODELS FOR FROZEN SOILS

Model classification based on its scale of representation:

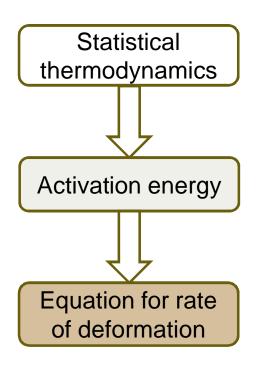
- Microscopic view
- > The theory of rate process
- Damage creep model

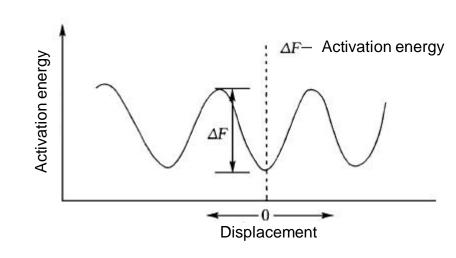
Phenomenological view

- Empirical model
- > Time Hardening Theory
- Elementary element based creep model

CREEP MODELS FROM MICROSCOPIC VIEW_1

> The theory of rate process: View deformation as a thermal activation process





$$\dot{\varepsilon} = \bar{C} \frac{kT}{h} \exp\left(-\frac{E}{RT}\right) \exp\frac{\Delta S}{k} \left(\frac{\sigma}{\sigma_0}\right)^{n+m}$$

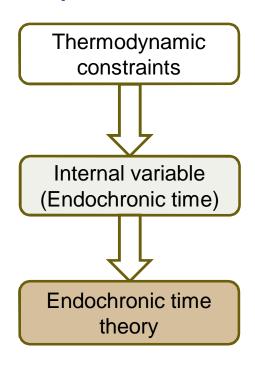
<u>from Fish 1983</u>

Related studies: Andersland (1967), Assur (1980)

A reasonable physical description Difficulty: Parameters obtained qualitatively by current testing technology

CREEP MODELS FROM MICROSCOPIC VIEW 2

Endochronic time theory: irreversible thermodynamic process of dissipative material



Constitutive relationship from Gopal et al. 1985

Deviatoric: z
$$de_{ij} = de_{ij}^e + de_{ij}^p$$

Volumetric:z' $d\varepsilon = d\varepsilon^e + d\varepsilon^p$

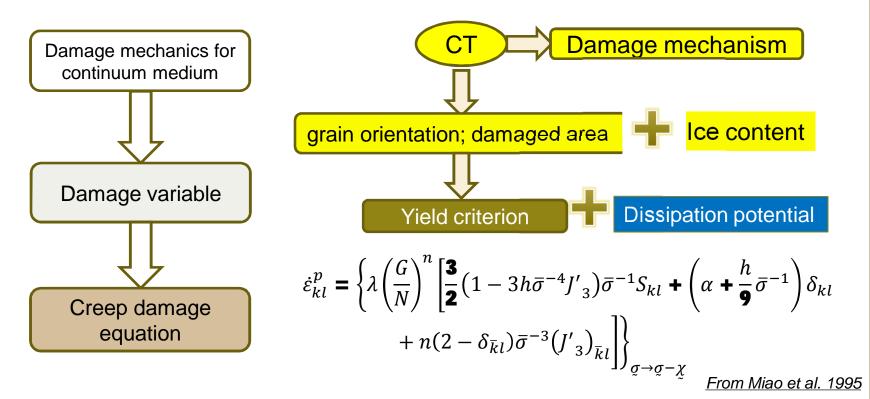
Definition $z \text{ and } z'$ $(dz')^2 = \left(\frac{d\varsigma}{z_1}\right)^2 + \left(\frac{dt}{\tau_1}\right)^2$
 $(dz')^2 = \left(\frac{d\varsigma'}{z_2}\right)^2 + \left(\frac{dt}{\tau_1}\right)^2$

Parameters determined by Bazant et al. 1983

- 1) Strain hardening and softening law;
- 2) Expansion and contraction
- 3) Hydrostatic pressure sensibility

CREEP MODELS FROM MICROSCOPIC VIEW 2

> The damage creep model: damage or recovery of soil structure



Following thermodynamics; Unique internal variable in frozen soil: ice

Contentry: Calibration of Parameters for thermodynamics and damage mechanics

CREEP MODELS FROM PHENOMENOLOGICAL VIEW_1

Empirical methods: a mathematical description of creep curve

Basic form:

Classified by creep stages:

- 1 Stress-strain-time
- 2 Stress-strain rate-time
- Primary creep model (Vyalov, 1966)
- II. Secondary creep model (Ladanyi, 1972)
- III. Tertiary creep model (?)

Simple structure; conveniently applied in simple engineering analysis (first estimation)

1) Poor versatility; 2) do not reflect internal mechanism

CREEP MODELS FROM PHENOMENOLOGICAL VIEW 2

> Time hardening model

Stress-strain-time:
$$\gamma_i^{cr} = f(\tau, \sigma_m) \phi(t) \omega(\theta)$$

Klein and Jessberger (1979):
$$\varepsilon = \sigma / E_0 + A\sigma^b t^c$$

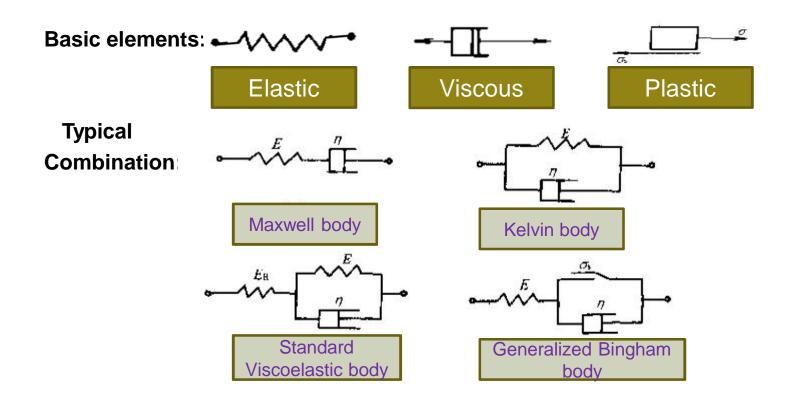
Herzog and Hofer (1981):
$$\varepsilon = \sigma^a t^b + dt$$

Simple; direct

Only available for constant temperature

CREEP MODELS FROM PHENOMENOLOGICAL VIEW 3

Elementary creep model: combination of mechanical elements



SUMMARIZATION

- Some are too complicated in form, too many parameters, even impossible to be obtained from conventional tests
- Some are lacking mechanism, just mathematical description
- Some are difficult to accommodate different thermal or load conditions

We need something new.

AGENDA

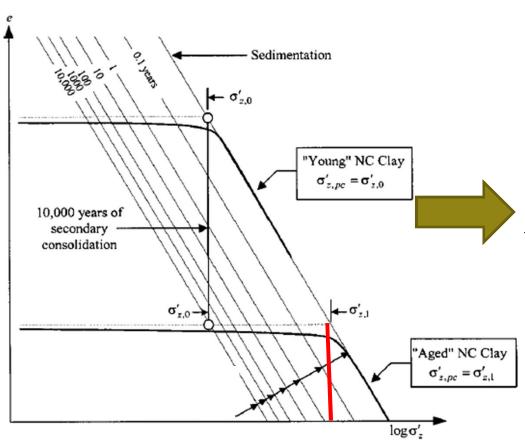
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BORROW THEORIES FROM UNFROZEN SOILS

• Ladanyi (1999): creep of frozen soils is not very much different from that of unfrozen soils

• Warm frozen soil is between frozen and unfrozen soils, most likely closer to unfrozen soils

-Creep of unfrozen soils based on p_c



Yin, et al. (1989):

$$\dot{\varepsilon}_{z} = \frac{k/V}{\sigma_{z}'} \sigma_{z}' + \frac{\psi/V}{t_{0}} \exp \left[-\left(\varepsilon_{z} - \varepsilon_{z0}^{ep}\right) \frac{V}{\psi} \right] \left(\frac{\sigma_{z}'}{\sigma_{p_{c}}'}\right)^{\lambda/\psi}$$

Vermeer and Neher(1999):

$$\varepsilon = \varepsilon^{e} + \varepsilon^{c} = A \frac{\sigma'}{\sigma'} + \frac{C}{\tau} \left(\frac{\sigma'}{\sigma_{p}} \right)^{\frac{E}{C}}$$

$$\sigma_p = \sigma_{p_c} \exp\left(\frac{\varepsilon^c}{B}\right)$$

(After Bjerrum, 1973)

Questions

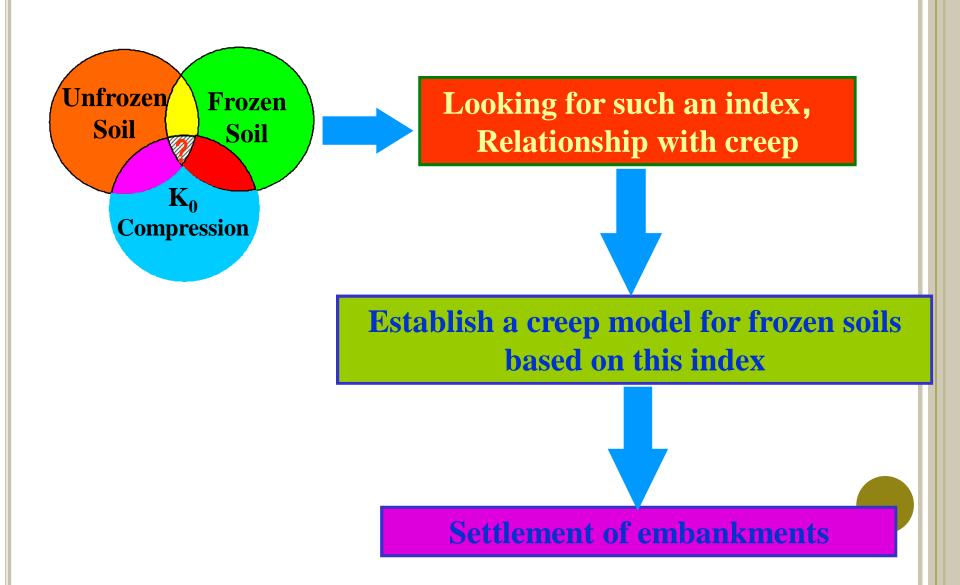
Is there an index in frozen soils similar to p_c ?

If so

What is its relationship with creep?

How is it possible to apply creep model of unfrozen soils to frozen soils?

Methodology



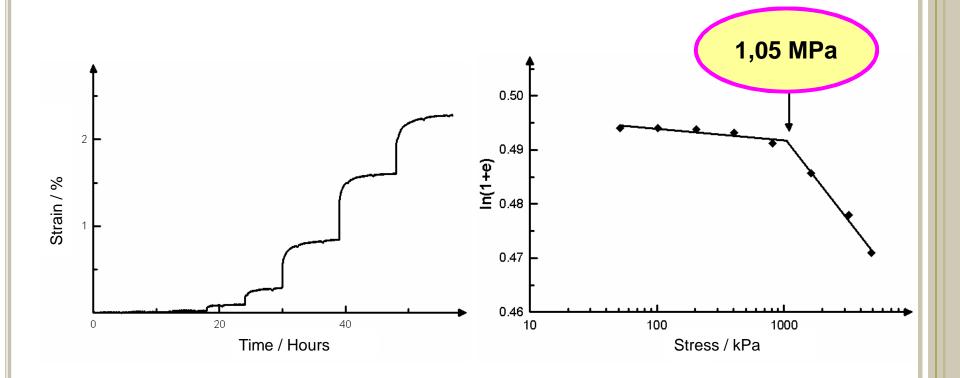
Testing

Phases	Purpose	Conditions	Samples
1	Prove the existence of an index similar to Pc	Same T Different γ _d	4
		Same γ _d Different T	4
2	Influence of creep on this index	γ _d , T, preload	30
3	Comprehensive analysis Relationship: Creep vs. Pc	Orthogonal design: γ _d , T, preload, creep time	10

48 creep tests so far

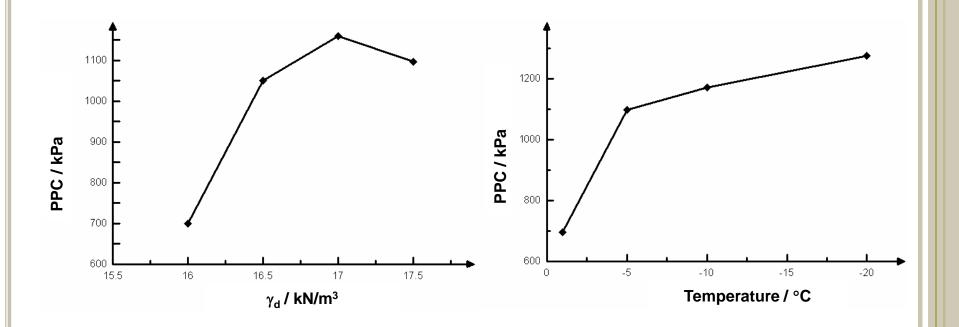
K₀ Test

Pseudo preconsolidation pressure PPC for frozen soils



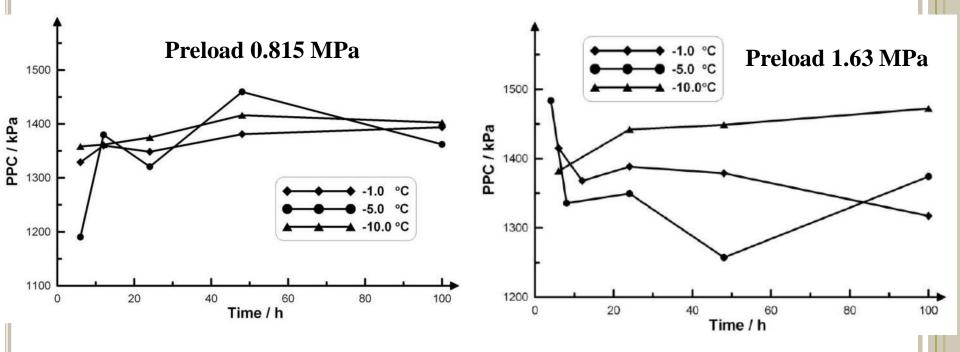
In " $\ln(1+e)$ " coordinates, there is clearly such an index

PPC: Mechanical behavior of frozen soils



- \triangleright PPC increases with the increase in γ_d , then does not change obviously
- >PPC increases linearly with the decrease in temperature

Well reflects the bonding in frozen soils



- > When preload is less than original PPC, PPC increases with time
- > When preload is larger than original PPC, PPC decreases with time

We are not ready to get a relationship between PPC and temperature, creep time; but we successfully proved the existence of such an index.



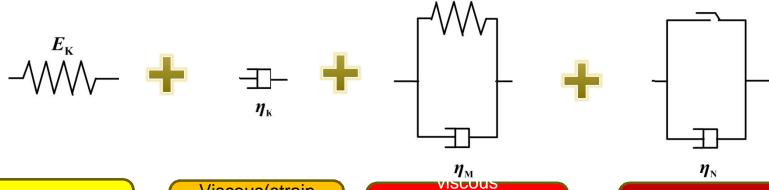
CREEP

SOIL

(Dr. Songhe Wang)

 $\phi\langle F\rangle$

> Establishing the model



Instantaneous elastic

Viscous(strain rate approaches zero)

(strain rate increases with

Viscoplastic with a yield surface

The creep model is

$$e_{ij} = \frac{S_{ij}}{2G_M} + \frac{S_{ij}}{2H_M}t + \frac{S_{ij}}{2G_K} \left[1 - \exp\left(-\frac{G_K}{H_K}t\right) \right] \quad \left(\phi(F) \le 0\right)$$

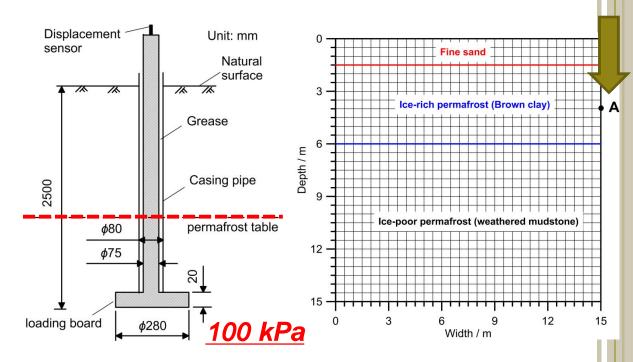
$$e_{ij} = \frac{S_{ij}}{2G_M} + \frac{S_{ij}}{2H_M}t + \frac{S_{ij}}{2G_K} \left[1 - \exp\left(-\frac{G_K}{H_K}t\right) \right] + \frac{1}{2H_N} \left\langle \phi(F) \right\rangle \frac{\partial Q}{\partial \{\sigma\}}t \quad \left(\phi(F) > 0\right)$$

Yield criterion(Ma, et al. 1994)
$$F = \sqrt{3J_2} - c - \sigma_m \tan \varphi + \frac{\tan \varphi}{2p_m} \sigma_m^2$$

> Long-term load test

Field load test

Loading pile

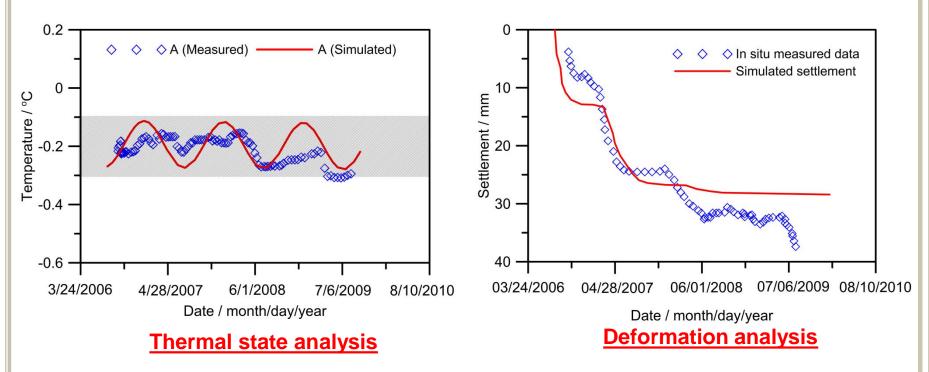


Numerical model

Only creep of underlying permafrost was considered

> Numerical simulation

After implementation of this model,



A simple model for creep of frozen soil might provide a way in engineering analysis.

A VISCO-HYPOPLASTIC CONSTITUTIVE MODEL FOR FROZEN SOIL

(Dr. Guofang Xu, Prof. Wei Wu)

$$\dot{\boldsymbol{\sigma}} = \dot{\boldsymbol{\sigma}}_{\mathrm{s}} + \dot{\boldsymbol{\sigma}}_{\mathrm{d}}$$

 $\sigma_{\rm s}$ is static stress, $\sigma_{\rm d}$ is dynamic stress.

Static part

$$\dot{\boldsymbol{\sigma}}_{s} = c_{1} [\operatorname{tr}(\boldsymbol{\sigma}_{s} - \boldsymbol{c})] \dot{\boldsymbol{\varepsilon}} + c_{2} \frac{\operatorname{tr}[(\boldsymbol{\sigma}_{s} - \boldsymbol{c}) \dot{\boldsymbol{\varepsilon}}]}{\operatorname{tr}(\boldsymbol{\sigma}_{s} - \boldsymbol{c})} (\boldsymbol{\sigma}_{s} - \boldsymbol{c}) + f_{\varepsilon} \cdot f_{cd} \cdot \left[c_{3} (\boldsymbol{\sigma}_{s} - \boldsymbol{c})^{2} + c_{4} (\boldsymbol{\sigma}_{s} - \boldsymbol{c})_{d}^{2} \right] \frac{\|\dot{\boldsymbol{\varepsilon}}\|}{\operatorname{tr}(\boldsymbol{\sigma}_{s} - \boldsymbol{c})}$$

in which c is the cohesion of frozen soil, f_{ε} is a scalar function of deformation, $f_{\varepsilon d}$ is a factor of creep damage.

$$f_{\varepsilon} = 2 - \exp(\alpha \cdot l + \beta)$$

 α and β are parameters, l is the accumulation of deformation.

$$l = \int_{t_0}^t \|\dot{\mathbf{\varepsilon}}(\tau)\| \,\mathrm{d}\tau$$

$$f_{\rm cd} = 1 + \gamma \cdot \int_{t_1}^{t_2} \langle \ddot{\varepsilon}(\tau) \rangle d\tau$$

 γ is a parameter, $\langle \ \rangle$ is Macaulay brackets.

Dynamic part

$$\dot{\boldsymbol{\sigma}}_{d} = \eta_{1} \sqrt{\eta_{2}^{2} + \operatorname{tr}(\dot{\boldsymbol{\varepsilon}}^{2})} \ddot{\boldsymbol{\varepsilon}}$$

 η_1 and η_2 are parameters, $\ddot{\epsilon}$ is strain acceleration.

Complete constitutive model - Rate dependent

$$\dot{\boldsymbol{\sigma}} = c_1 [\operatorname{tr}(\boldsymbol{\sigma} - \boldsymbol{c})] \dot{\boldsymbol{\varepsilon}} + c_2 \frac{\operatorname{tr}[(\boldsymbol{\sigma} - \boldsymbol{c}) \dot{\boldsymbol{\varepsilon}}]}{\operatorname{tr}(\boldsymbol{\sigma} - \boldsymbol{c})} (\boldsymbol{\sigma} - \boldsymbol{c}) + f_{\varepsilon} \cdot f_{\operatorname{cd}} \cdot \left[c_3 (\boldsymbol{\sigma} - \boldsymbol{c})^2 + c_4 (\boldsymbol{\sigma} - \boldsymbol{c})_{\operatorname{d}}^2 \right] \frac{\|\dot{\boldsymbol{\varepsilon}}\|}{\operatorname{tr}(\boldsymbol{\sigma} - \boldsymbol{c})} + \eta_1 \sqrt{\eta_2^2 + \operatorname{tr}(\dot{\boldsymbol{\varepsilon}}^2)} \, \dot{\boldsymbol{\varepsilon}}$$

> CALIBRATION OF THE CONSTITUTIVE MODEL

Parameters in the static part

When the two linear and nonlinear terms in the static part of the model are abbreviated as L_1 , L_2 , N_1 and N_2 , this part can be rewritten as:

$$\dot{\boldsymbol{\sigma}}_{s} = c_{1} \mathsf{L}_{1}(\boldsymbol{\sigma}_{s}) : \dot{\boldsymbol{\varepsilon}} + c_{2} \mathsf{L}_{2}(\boldsymbol{\sigma}_{s}) : \dot{\boldsymbol{\varepsilon}} + c_{3} \mathsf{N}_{1}(\boldsymbol{\sigma}_{s}) \|\dot{\boldsymbol{\varepsilon}}\| + c_{4} \mathsf{N}_{2}(\boldsymbol{\sigma}_{s}) \|\dot{\boldsymbol{\varepsilon}}\|$$

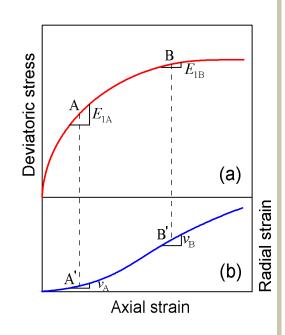
In a conventional triaxial test, owing to $\dot{\sigma}_2 = \dot{\sigma}_3 = 0$, the above equation can be divided into two scalar equations as follows:

$$\dot{\sigma}_1 = c_1 L_{11} \dot{\varepsilon}_1 + c_2 L_{12} \dot{\varepsilon}_3 + c_3 N_{11} \sqrt{\dot{\varepsilon}_1^2 + 2\dot{\varepsilon}_3^2} + c_4 N_{12} \sqrt{\dot{\varepsilon}_1^2 + 2\dot{\varepsilon}_3^2}$$

$$\dot{\sigma}_3 = c_1 L_{21} \dot{\varepsilon}_1 + c_2 L_{22} \dot{\varepsilon}_3 + c_3 N_{21} \sqrt{\dot{\varepsilon}_1^2 + 2\dot{\varepsilon}_3^2} + c_4 N_{22} \sqrt{\dot{\varepsilon}_1^2 + 2\dot{\varepsilon}_3^2}$$

Owing to the radial stiffness $E_{A3} = E_{B3} = 0$, we have

$$\begin{split} E_{\rm A1} &= c_1 L_{11} + c_2 L_{12} v_{\rm A} + c_3 N_{11} \sqrt{1 + 2 v_{\rm A}^2} + c_4 N_{12} \sqrt{1 + 2 v_{\rm A}^2} \\ 0 &= c_1 L_{21} + c_2 L_{22} v_{\rm A} + c_3 N_{21} \sqrt{1 + 2 v_{\rm A}^2} + c_4 N_{22} \sqrt{1 + 2 v_{\rm A}^2} \\ E_{\rm B1} &= c_1 L_{11} + c_2 L_{12} v_{\rm B} + c_3 N_{11} \sqrt{1 + 2 v_{\rm B}^2} + c_4 N_{12} \sqrt{1 + 2 v_{\rm B}^2} \\ 0 &= c_1 L_{21} + c_2 L_{22} v_{\rm B} + c_3 N_{21} \sqrt{1 + 2 v_{\rm B}^2} + c_4 N_{22} \sqrt{1 + 2 v_{\rm B}^2} \end{split}$$



The parameters c_i (i = 1, ..., 4) can be obtained by solving the above equation system with respect to the variables c_i .

Parameters α and β

Parameters α and β can be obtained from the following expressions:

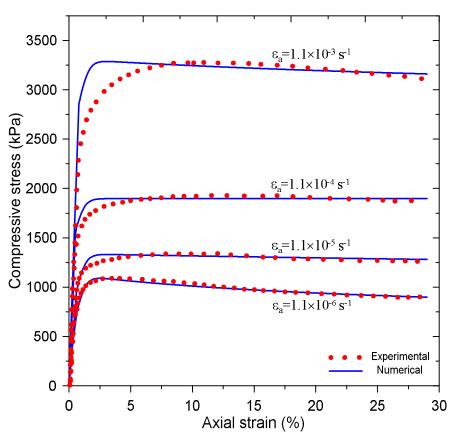
$$\alpha = 1 - \left(T/T_{\text{ref}}\right)^{n_1}$$
$$\beta = -\left(T/T_{\text{ref}}\right)^{n_2}$$

$$\beta = -(T/T_{\rm ref})^{n_2}$$

in which T_{ref} is a reference temperature and can be regarded as -1°C (Zhu and Carbee, 1984).

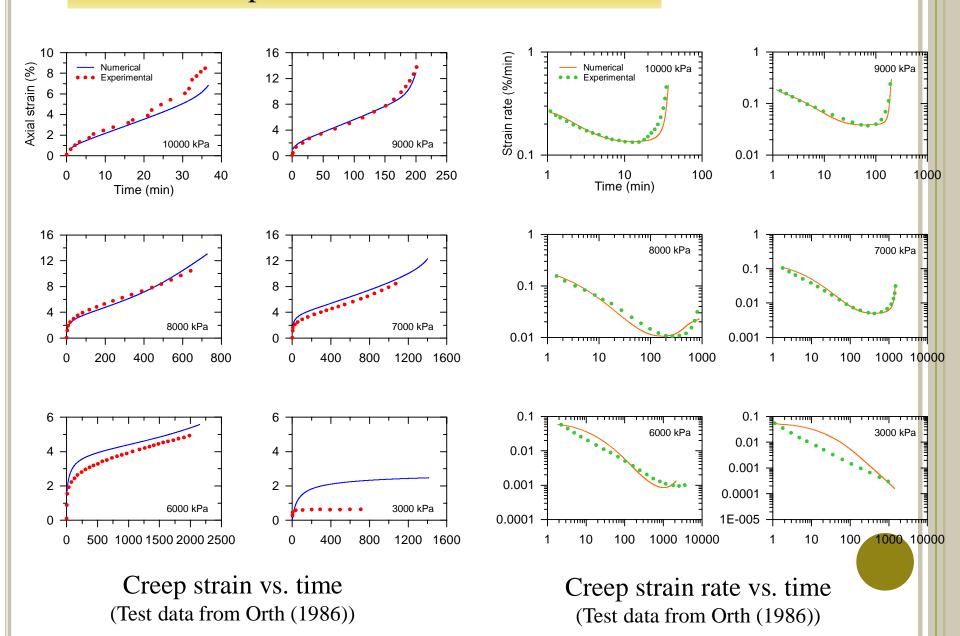
Parameters η_1 and η_2 in the dynamical part can only be obtained by fitting the experimental data, as done by Hanes and Inman (1985).

➤ VERIFICATION OF THE CONSTITUTIVE MODEL Uniaxial compression tests at different loading rates



Stress-strain relationship at different strain rates (Data from Zhu and Carbee, 1984)

Uniaxial creep tests at different stress levels



CONCLUDING REMARKS

- General features in stress-strain-time curves for frozen are similar to that of unfrozen soils. Warm frozen soil is closer to unfrozen soils.
- A warm frozen soil is defined according to mechanical properties. Its creep was successfully observed in situ.
- No generally recognized constitutive models are found for creep of frozen soils. We have tried in different ways.

ACKNOWLEDGEMENTS

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